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Article

Effects of Nonaerated Circulation Water Velocity on Nutrient Release from Aquaculture Pond Sediments

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Abstract: Sustaining good water quality in aquaculture ponds is vital. Without an aerator, the dissolved oxygen in ponds comes primarily from mass transfer at the water–ambient atmosphere interface. As sediment can seriously affect water quality, this study used indoor experiments to examine the nutrient (nitrogen and phosphorus) release mechanisms and fluxes from sediment in aquaculture ponds with moving water but no aeration. The results showed that the ammonia nitrogen (NH₃-N) concentration in the overlying water was inversely proportional to flow velocity and that a higher flow velocity tended to result in a lower concentration in the overlying water, a steeper vertical gradient of concentration within the bed sediments, and a faster release rate from the sediments. The sediment disturbed by flowing water released more nitrate nitrogen (NO₃-N) and nitrite nitrogen (NO₂-N) into the overlying water and NO₂-N could become oxidized into NO₃-N. In still water, NO₃-N was released gradually and some anaerobic NO₃-N was nitrified into NO₂-N. Phosphorus release from the sediments was controlled by the adsorption–desorption balance, with the phosphorus concentration in the overlying water dropping gradually to a steady value from its initial maximum. The relationship between NH₃-N release flux and flow rate is described by a cubic function.

Keywords: nutrient release; sediments; circulation water; aquaculture

1. Introduction

The rapid development of the aquaculture industry in recent years has led to abundant residual feed, metabolites, and excreta being deposited in pond sediments. The resulting 50- to 500-fold rise in the nitrogen (N) and phosphorus (P) concentrations in the sediments compared to those in the overlying water [1] vastly increases the likelihood that N and P will be released from the sediment–water interface and migrate into the water, potentially threatening the aquaculture environment and impacting aquaculture production. Given that the N and P released can also boost autotrophic growth in freshwater ecosystems, seriously affecting the biomass and population structures of aquatic species [2–4], a great deal of research has focused on the role played by N and P in aquaculture [5,6]. Most previous studies examined the changes in nutrition concentration in the sediment porewater and overlying water and measured the nutrient diffusion flux under still water conditions [7], but the water turbulence caused by storms, aeration, and bioturbation can have a significant impact on the release of N and P at the sediment–water interface [8]. The deposition and resuspension of sediments resulting from water turbulence can prompt the release of nutrients and metals into overlying water [9]

and increase the exchange coefficient between the overlying water and sediments, thus accelerating the diffusion of any contaminants present in the sediments into the overlying water and boosting nutrient release. Reddy et al. [8] found that in Apoka Lake, which is located in the US state of Florida, under turbulent water conditions the nutrient concentrations in the overlying water resulting from the deposition and resuspension was dozens of times higher than those resulting from diffusion alone. In Denmark's Arreso Lake, Sondergaard et al. [10] found that sediment resuspension led to 20- to 30-fold increases in the nutrient concentrations in the overlying water under similar conditions.

The dissolved oxygen (DO) concentration in water also influences the release of N and P at the sediment-water interface. Usui et al. [11] used ^{15}N as a tracer to monitor the fate of nutrients in sediments in the Tama estuary, Japan. By comparing changes in the level of N_2O , they showed that increasing the concentrations of NO_3^- and DO in the overlying water stimulated nitrification and denitrification. In a series of laboratory experiments, Baudo [12] demonstrated that while insoluble ferric hydroxide ($\text{Fe}(\text{OH})_3$) was converted into soluble $\text{Fe}(\text{OH})_3$, there was a massive release of iron bound phosphorus (Fe-P) into the water column under anaerobic conditions, one of the potential mechanisms leading to the eutrophication of shallow lakes. Although phosphorus release can occur under aerobic conditions, the amount released is generally thought to be much smaller than under anaerobic conditions. The aerobic release mechanism is primarily based on the mineralization of sediments and the aerobic decomposition of organic matter.

The objective of this study was therefore to determine precisely how the N and P release mechanisms in aquaculture sediments are influenced by water hydrodynamics and DO conditions. To this end, a series of experiments were conducted to simulate circular flow in the overlying water and measure the N and P released from sediments in an aquaculture pond that was not subject to aeration.

2. Experimental Apparatus and Method

2.1. Experimental Apparatus

The experimental apparatus is illustrated in Figure 1. The tank, made of clear poly(methyl-methacrylate) organic glass, was 50 cm long, 25 cm wide, and 35 cm high. It was wrapped in aluminum foil to deter algae growth and reproduction under light stimulation. The tank cover was equipped with a small hole to provide access for the DO meter probe and for water sampling; this was sealed with a rubber plug when not in use. Two vertical 10-cm-high glass plates were placed inside the tank, 5 cm away from the water inlet and outlet pipes, as shown in Figure 1. The sediments were placed between these two plates to protect the sediment surface from being washed away by the water flow. Three small 6-mm-diameter holes were drilled along one side of the tank, with the first being about 1 cm above the sediment surface and the others 3 cm and 6 cm, respectively, below the sediment surface, and rhizosphere soil solution collectors (Rhizon pipes) inserted horizontally into each to collect porewater samples. The structure of the Rhizon pipes is illustrated in Figure 2. The sample sediments for the laboratory experiments were taken from an aquaculture pond, containing three species of fish, namely black, grass, and silver carp, with a breeding density of about $6000 \text{ kg}\cdot\text{ha}^{-1}$. These sediments consisted of muddy fine sandy soils with a mean particle size of about 0.1 mm. Distilled water was used as the overlying water to eliminate any effects of algae and microorganism in natural water. In order to conduct synchronous measurements under different flow rates, five identical sets of the apparatus shown in Figures 1 and 2 were constructed, each with its own pump, to allow the circulation flow rates to be adjusted independently. Seen from Figure 1, the pipes, the valve, water pump, flow meter, and tank were composed of a water flow process. Along with the arrow direction, water in the tank flowed circularly, with the water pump adjusting the water flow rate.

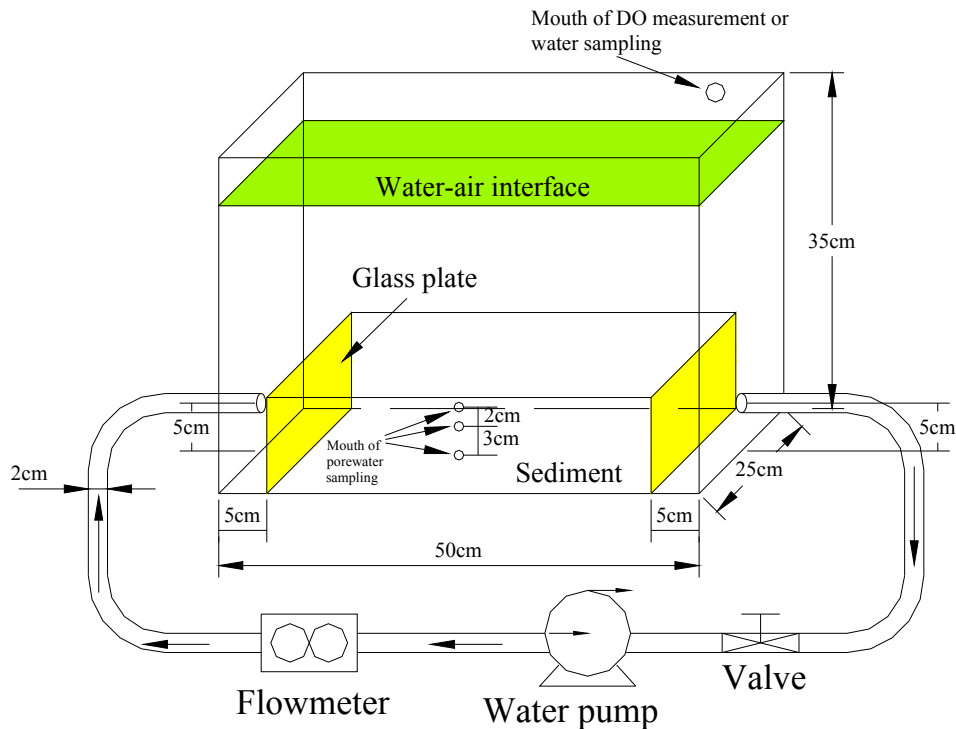


Figure 1. Illustration of the experimental apparatus.

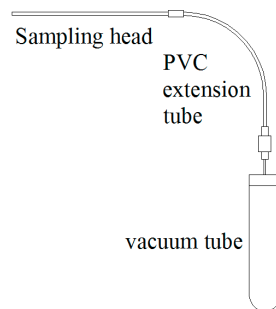


Figure 2. Structure of the Rhizon pipes.

The geochemical processes of porewater govern important aspects of all lake and oceanic water environments, most notably the spatial distribution, migration, and transformation of chemical substances at the sediment-water interface as well as the early diagenetic process within sediments [13]. As part of the effort to develop a better understanding of how these processes function, some researchers have used balance sampling devices to collect in-situ porewater, while others have used specially designed sediment collectors to harvest clay cores or superficial sediments that are then layered, centrifuged or filter pressed to extract the porewater [14,15]. For laboratory experiments, porewater is often acquired via an anaerobic centrifugation approach [8,16,17], but this approach relies on strict operational procedures and can only be conducted in a nitrogenous environment to avoid oxidization effects. The chemical balance of the porewater is also all too easily influenced by factors such as the pH and temperature. The biggest shortcoming of the anaerobic centrifugation approach, however, is the resulting destruction of the biologic characteristics of porewater, making it unsuitable for long-term sampling. Giesler et al. [13] used a high-speed centrifugal method and a zero-tension lysimeter to measure the solute leaching loss, and found that the leaching loss was overestimated substantially. In contrast, measurements made using the type of Rhizon pipe utilized in this study, which has a 10-cm-long sampling head and a 12-cm-long PVC (polymer of vinyl chloride) extension

tube, proved more effective than the lysimeter for collecting porewater samples within sediments because it minimizes sediment disturbance and changes in the porewater properties, isolates the sample from air, and is both inexpensive and reusable [18]. The head incorporates a negative-pressure joint and is connected to a vacuum tube in which the porewater sample is collected and stored (Figure 2).

2.2. Experimental Method

The sediment sample was collected in the summer of 2014 from an aquaculture farm in Guangzhou, China using a clamshell sampler and immediately sealed in a polyethylene plastic bucket for transport back to the laboratory. Upon arrival, large stones, organic detritus, benthos, and rubbish were removed, after which the sample was stirred and mixed evenly to form a rebonded sediment sample with average physicochemical properties for the sediments in that aquaculture pond, which is typical of those in the area. The space between the two glass plates in each tank was filled with rebonded sediment to a depth of 10 cm, level with the top of the glass retaining plates, and the sediment surface smoothed gently. During this process, the three Rhizon pipes were inserted at the appropriate depths in each tank. After the sediment had been allowed to settle for 15 days, distilled water was slowly poured into each tank until the overlying water depth reached about 18 cm.

The sediments were cultured for 7 days under different circulation flow rates. During this period, porewater samples from within the sediments and in the overlying water were collected using vacuum pipes at the same time every day. The resulting 5 mL samples were then used to test ammonia nitrogen ($\text{NH}_3\text{-N}$), nitrite nitrogen (NO_2^-), nitrate nitrogen (NO_3^-), and soluble reactive phosphorus (SRP) via an enzyme linked immunosorbent assay (ELISA). Due to its strong specificity, high sensitivity, good repeatability, simple operation, and freedom from radioactive contamination [19], ELISA is capable of testing hundreds of samples very quickly, greatly facilitating the collection and analysis of this type of data. The pH of the overlying water was measured by a pH meter (Acme Meter Corp, Springfield, IL, USA), while overlying water temperatures (T) and DO concentration were measured by a YSI ProODO optical dissolved oxygen instrument (YSI Inc., Yellow Springs, OH, USA). Oxidation-reduction potentials (ORP) of the overlying water were tested by a Model oxidization-reduction potential meter.

To determine the mean flow velocity in the overlying water, the mean water depth in each of the experimental tanks was calculated as:

$$h = \frac{V - 2lWh_p}{LW} \quad (1)$$

where h is the mean overlying water depth; V is the water volume of the tank (= 27.5 L); W is the width of the tank (= 25 cm); L is the length of the tank (= 50 cm); l is the distance from the plate to the nearest side wall of the tank (= 5 cm); and h_p is the height of the glass plate (= 10 cm).

The mean flow velocity in the overlying water was calculated as:

$$v = \frac{Q}{hW} \quad (2)$$

where v is the mean flow velocity in the overlying water; and Q is the flow rate in the pipe.

Based on Equations (1) and (2), the mean overlying water depth h was calculated to be 0.2 m and the mean flow velocities in the overlying water in the five tanks were calculated to be 0, 0.47, 0.54, 0.57, and 0.61 $\text{cm}\cdot\text{s}^{-1}$, respectively.

3. Results

As stated above, the release of N and P from sediments can be affected by a number of external factors [20], including organic matter supplements [21–23]; temperature, DO, pH, and ORP [24]; and the physical properties of the sediments [25]. We will therefore analyze these factors before moving

on to discuss the effect of the flow velocity of the overlying water on the N and P release rates from the sediments.

3.1. Changes in the External Factors Affecting the Overlying Water during Culture Periods

3.1.1. Water Temperature

Water circulation driven by pumps inevitably emits some heat, slightly raising the water temperature. In our experiments, water temperature was measured daily and the results are shown in Figure 3. The water temperature in the still-water tank (i.e., with no pumping) was lower, as it was only influenced by air temperature, but the water temperatures in the other four tanks ($v = 0.47, 0.54, 0.57,$ and $0.61 \text{ cm}\cdot\text{s}^{-1}$) were very similar. Water temperature is known to indirectly influence P release in sediments [26], with a high temperature boosting the activity of some substances and accelerating physical, chemical, and biological reactions such as the diffusion and degradation of organic matter. Microorganisms can also decompose organic P in sediments into inorganic P and convert insoluble phosphorus compounds into soluble ones and higher metabolic rates can accelerate the degradation of organic matter [27], thus promoting P release from sediments. For N release, Liu et al. [28] reported that water temperature can influence degradation speed, when temperatures drop too low, the degradation of N in deeper-sediments can cease altogether. Based on a series of experiments on sediment contaminant release, Huang et al. [29] found that a high temperature accelerates the emission of dissolved contaminants considerably. Ogilvie et al. [30] agreed, pointing out that rising temperatures not only favor ammonification and nitrification but also enhance ammonifying and nitrifying bacterial activity.

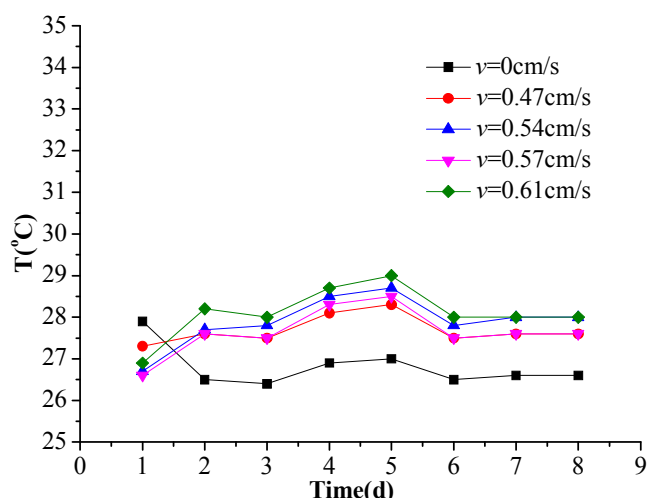


Figure 3. Plot of overlying water temperature for different flow rates during the experiment.

3.1.2. DO and ORP

The DO levels in the overlying water and the corresponding ORP values are presented in Figure 4a,b, respectively. ORP is a measure of the relative degree of water oxidizability or reducibility. Wetzel et al. [31] and Seiki et al. [32] showed that the DO content in the overlying water is an important factor influencing ORP. As the data in Figure 4a shows, even when the culture tanks were sealed and not aerated, the existence of an air-filled space above the overlying water provided some air re-aeration through water surface entrainment caused by the circulating water flow, as indicated by the declining concentrations of DO over the first two days and the subsequent increases later in the experiment, which contrast markedly with the DO concentrations in the still-water tank, which declined monotonically throughout. Interestingly, the data in Figure 4b show that the ORP was

consistently negative, ranging from -60 to -100 , implying that the culture water environment favors the reduction of ORP.

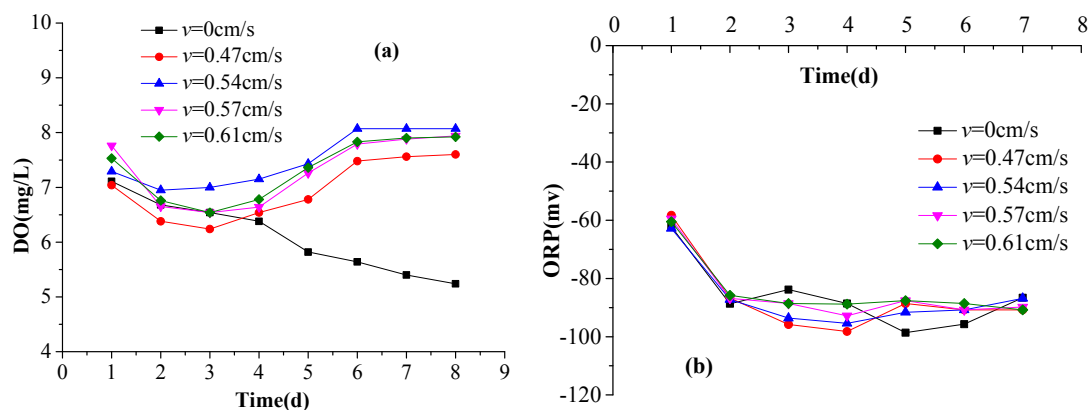


Figure 4. (a) Dissolved oxygen (DO) concentrations; and (b) oxidation-reduction potentials (ORP), in the overlying water during the experiment.

3.1.3. pH

For all the flow rates tested, the culture water pH was effectively independent of flow rate varying only from 6.5 to 7.0 (Figure 5), indicating a weak acidity, although pH did tend to decrease with culture time. The pH is known to have a significant impact on P morphology [33], however, as iron, aluminum, and cadmium in water and sediments are highly likely to react with soluble inorganic P in acidic solutions, forming non-dissolvable deposits, while anions of phosphate are more likely to be absorbed by bauxite, kaolinite, and/or montmorillonite, as well as by newly precipitated $\text{Fe}(\text{OH})_3$ and aluminum hydroxide ($\text{Al}(\text{OH})_3$) colloids. The deposition-dissolution and adsorption-dissolution processes have a major influence on the physiochemical dynamics (in this case, the solubility and biological cycle) of P compounds. pH also plays an important role in N release from sediments, with the optimum pH for sediment nitrification being about 7.5 [34]. If the pH of the water increases, the rate of nitrification rises, leading to higher levels of free ammonia nitrogen. However, when the pH exceeds 7.5, the nitrification rate becomes almost independent of the free ammonia nitrogen concentration. During this experiment, the pH in the overlying water was kept weakly acid, with a pH below the optimum value, thus limiting the activity of the enzymes in the nitrifying bacteria and reducing the nitrification activity.

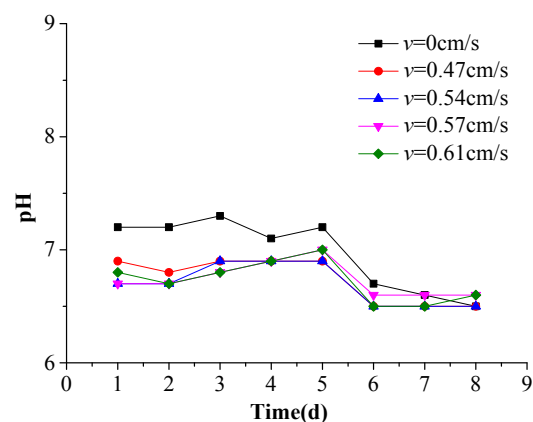


Figure 5. Variations in the pH of the overlying water during the experiment.

3.2. Effects of Flow Rate on NH₃-N Concentration

Figure 6 shows how the NH₃-N concentrations in the porewater and overlying water varied during the course of the experiment. For any given flow rate, the NH₃-N concentration near the surface of the sediment was lower than that at greater depths, while for a given depth, the NH₃-N concentration at a faster flow rate was lower than that at a slower flow rate. This is because the sediments were disturbed more by higher flow rates, thus releasing more NH₃-N into the overlying water and introducing more oxygen into the sediments to oxidize NH₃-N. The NH₃-N concentration gradient between the 6-cm- and 3-cm-deep sediments was steeper than the gradient between the 1-cm-deep sediment and the overlying water, demonstrating the effect of flow rate on NH₃-N release from the sediments. The gradient around the 3 cm depth was greatest because the sediments at this depth were both releasing NH₃-N upward and receiving DO from shallower sediments. In addition, the NH₃-N concentration in the 1-cm-deep sediments decreased gradually, ultimately achieving a steady state. This may be because the NH₃-N concentration gradient was becoming less steep, leading to a reduced release rate. Further, the NH₃-N concentration increased with time for any given flow condition, although a higher flow rate tended to be accompanied by a lower NH₃-N concentration. This is probably because higher flow rates would likely stir up the sediments more, thus encouraging the release of nutrients.

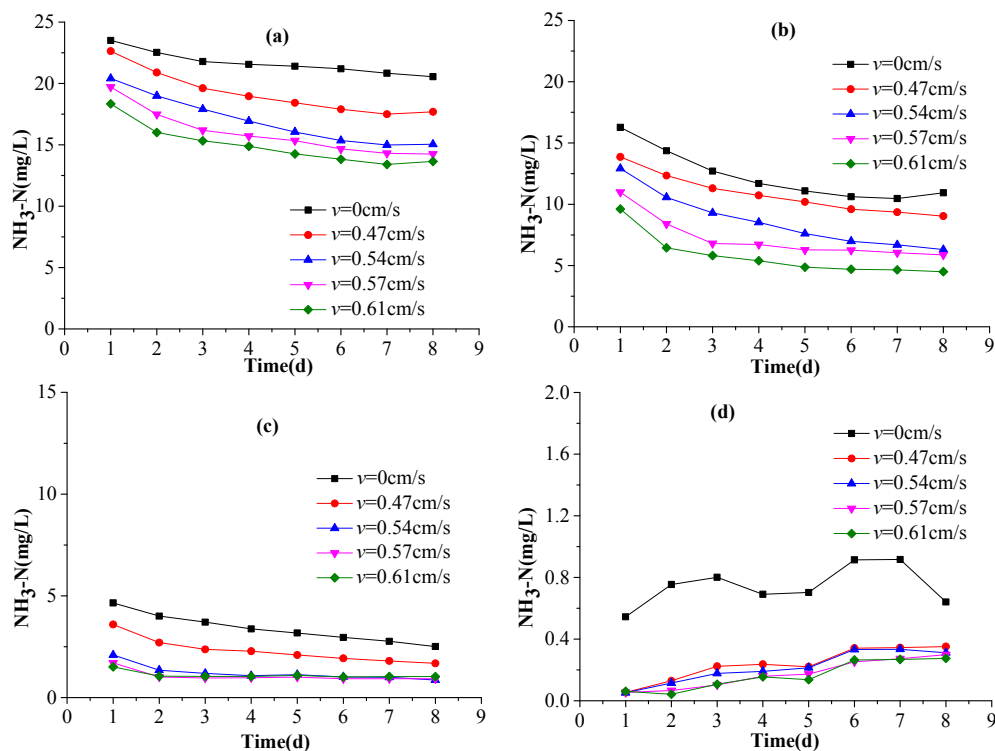


Figure 6. The concentrations of ammonia nitrogen (NH₃-N) in the sediment at depths of: (a) 6 cm; (b) 3 cm; (c) 1 cm; and (d) in the overlying water, throughout the experiment.

The release flux of NH₃-N from the sediments can be calculated as:

$$J = C(t) - C(t + 1) \quad (3)$$

where J is the sediment release flux ($\text{mg L}^{-1} \cdot \text{day}^{-1}$); and $C(t)$ and $C(t+1)$ are the measured NH₃-N concentrations in the sediment porewater on days t and $t + 1$, respectively.

The results obtained using Equation (3) are shown in Figure 7. The NH₃-N release fluxes at the three depths exhibited a similar pattern: NH₃-N was released quickly in the first 1–2 days, but then

the release slowed, gradually becoming stable. The initial rapid release is likely due to the greater NH₃-N concentration gradient between the overlying water (0 mg·L⁻¹) and the sediments; as time progresses, the NH₃-N concentration in the overlying water increases so the concentration gradient gradually decreases, leading to a slower NH₃-N release rate. For the higher flow rates, the sediments are disturbed more strongly, resulting in a higher initial release rate that is once again followed by a steady rate, which indicates the decreasing effect of flow rate on NH₃-N release. For the first few days, the flowing-water tanks released more NH₃-N than the still-water tank, but as time elapsed, the release flux in the flowing-water tanks tended to approach that in the still-water tank. Such decreasing trends in NH₃-N release rate can be well modeled using regression equations (Table 1), as indicated by the high values obtained for R², which ranged from 0.895 to 0.995.

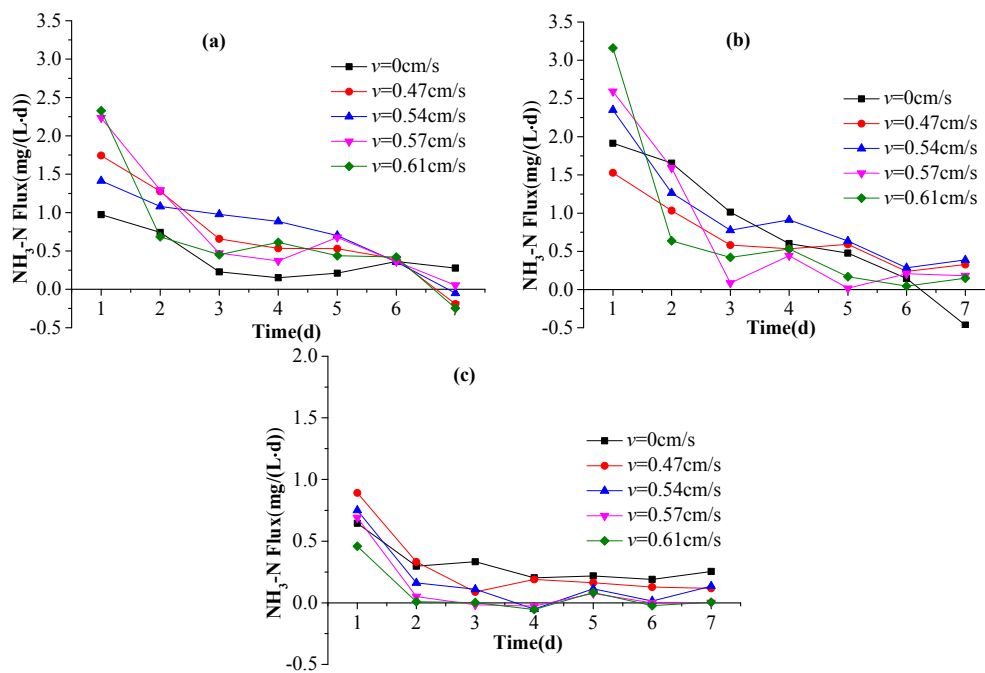


Figure 7. The release flux of NH₃-N during the course of the experiment measured at depths of: (a) 6 cm; (b) 3 cm; and (c) 1 cm within the sediments.

Table 1. The regression equations for the NH₃-N release flux.

Sediment Depth (cm)	Velocity <i>v</i> (cm·s ⁻¹)	Regression Equation (x: Day; y: Release Flux)	Coefficient of Determination R ²
6	<i>v</i> = 0 cm/s	$y = 1.75 - 0.858x + 0.152x^2 - 0.008x^3$	0.906
	<i>v</i> = 0.47 cm/s	$y = 3.026 - 1.526x + 0.333x^2 - 0.026x^3$	0.975
	<i>v</i> = 0.54 cm/s	$y = 1.853 - 0.58x + 0.134x^2 - 0.013x^3$	0.995
	<i>v</i> = 0.57 cm/s	$y = 4.337 - 2.57x + 0.562x^2 - 0.04x^3$	0.971
	<i>v</i> = 0.61 cm/s	$y = 4.903 - 3.415x + 0.829x^2 - 0.064x^3$	0.972
3	<i>v</i> = 0 cm/s	$y = 2.676 - 0.797x + 0.116x^2 - 0.009x^3$	0.982
	<i>v</i> = 0.47 cm/s	$y = 2.362 - 1.002x + 0.181x^2 - 0.012x^3$	0.952
	<i>v</i> = 0.54 cm/s	$y = 3.753 - 1.793x + 0.349x^2 - 0.023x^3$	0.949
	<i>v</i> = 0.57 cm/s	$y = 4.808 - 2.57x + 0.459x^2 - 0.027x^3$	0.937
	<i>v</i> = 0.61 cm/s	$y = 6.2 - 4.046x + 0.873x^2 - 0.06x^3$	0.921
1	<i>v</i> = 0 cm/s	$y = 0.979 - 0.439x + 0.08x^2 - 0.005x^3$	0.895
	<i>v</i> = 0.47 cm/s	$y = 1.802 - 1.167x + 0.257x^2 - 0.018x^3$	0.970
	<i>v</i> = 0.54 cm/s	$y = 1.519 - 0.991x + 0.204x^2 - 0.013x^3$	0.926
	<i>v</i> = 0.57 cm/s	$y = 1.622 - 1.229x + 0.284x^2 - 0.02x^3$	0.948
	<i>v</i> = 0.61 cm/s	$y = 1.089 - 0.838x + 0.195x^2 - 0.014x^3$	0.897

The generic format of these regression equations consists of a cubic function expressed as:

$$y = a + b_1x + b_2x^2 + b_3x^3 \quad (4)$$

where y is the $\text{NH}_3\text{-N}$ release flux from the sediments at day x ($\text{mg}\cdot\text{L}^{-1}\cdot\text{day}^{-1}$); and a , b_1 , b_2 , and b_3 are the regression coefficients.

A positive value for y indicates that $\text{NH}_3\text{-N}$ is released from the sediment into the overlying water, whereas a negative value means that $\text{NH}_3\text{-N}$ is absorbed from the overlying water back into the sediment. Because the overlying water in these experiments was distilled and free of nutrients, the sediments, which were rich in $\text{NH}_3\text{-N}$, initially released a large amount of $\text{NH}_3\text{-N}$ into the overlying water. However, once a balance in the $\text{NH}_3\text{-N}$ between the sediments and the overlying water had been achieved, the release from the sediments to the overlying water slowed or even ceased. Once this “balance point” was reached, the $\text{NH}_3\text{-N}$ in the overlying water began to be absorbed back into the sediments.

3.3. Effects of Flow Rate on $\text{NO}_3\text{-N}$ Concentration

Under flowing-water conditions, a lower flow rate corresponded to a higher $\text{NO}_3\text{-N}$ concentration for any given sediment depth, while a lower depth had a lower $\text{NO}_3\text{-N}$ concentration for the same test day (Figure 8). The $\text{NO}_3\text{-N}$ concentrations in the overlying water and at a sediment depth of 1 cm increased with time, whereas the $\text{NO}_3\text{-N}$ concentrations 3 cm and 6 cm below the surface of the sediments decreased after the initial increase. Under the still-water condition, the $\text{NO}_3\text{-N}$ concentration at a depth of 6 cm decreased monotonically with time.

One possible explanation for this observation is that because the sediment was disturbed by the water flow, more $\text{NO}_3\text{-N}$ and $\text{NO}_2\text{-N}$ was released from the sediments into the overlying water and $\text{NO}_2\text{-N}$ was more likely to be oxidized into $\text{NO}_3\text{-N}$. This is also probably true for the 1-cm-deep sediments, which were soaked by the overlying water. As time progressed, the $\text{NO}_3\text{-N}$ concentration in the overlying water and in the 1-cm-deep sediments increased, reducing the concentration gradient between the sediments and the overlying water. Comparing the DO variations (Figure 4a) with the $\text{NO}_3\text{-N}$ variations (Figure 8), a lower DO generally corresponded to a higher $\text{NO}_3\text{-N}$ concentration regardless of the flow rate or sediment depth, indicating that some $\text{NO}_2\text{-N}$ had probably been gradually oxidized into $\text{NO}_3\text{-N}$ due to aeration from the overlying water. This transformation process is supported by the variations in the $\text{NH}_3\text{-N}$ concentration (Figure 6). The $\text{NO}_3\text{-N}$ concentrations in the 3-cm- and 6-cm-deep sediments exhibited flipped “V-shaped” variation trends for the test flow rates, possibly as a result of the dynamic exchange of $\text{NO}_3\text{-N}$ between the upper and lower sediment levels. At the beginning of the experiment, the deep sediments released $\text{NO}_3\text{-N}$, leading to a decreased concentration and an increasing release rate.

For the still-water tank, because there was no flow turbulence $\text{NO}_3\text{-N}$ was likely to have continued to be released gradually throughout with some anaerobic $\text{NO}_3\text{-N}$ being nitrified into $\text{NO}_2\text{-N}$, leading to a monotonic decreasing trend for $\text{NO}_3\text{-N}$ (Figure 8). Chong [35] found that the nitrate release rate in deep water is slower than that in shallow water, which implies that the nutrient cycle in deep water is slower because there is less vertical mixing. This finding is supported by our experimental results, which confirmed that stronger vertical mixing in shallower water can induce more nutrient release. The overlying water temperature was relatively higher on the 4th and 5th days (Figure 3), while the $\text{NO}_3\text{-N}$ concentrations were somewhat reduced on those days (Figure 8b–d). In contrast, from the 6th day onward, the overlying water temperature decreased but the $\text{NO}_3\text{-N}$ concentrations started to increase. This indicates that the $\text{NO}_3\text{-N}$ concentration in the sediments is likely negatively related to water temperature, which is consistent with the findings reported by previous studies. For example, Anthony and Lewis [36] examined the release of endogenous nitrogen in an oligotrophic lake, and found that the release rate of $\text{NH}_3\text{-N}$ at the sediment-water interface was positively correlated with water temperature while the $\text{NO}_3\text{-N}$ concentration in the overlying water was negatively correlated

with water temperature. The NO₃-N variations with time can once again be expressed as regression equations (Table 2), with high values of R², ranging from 0.769 to 0.973.

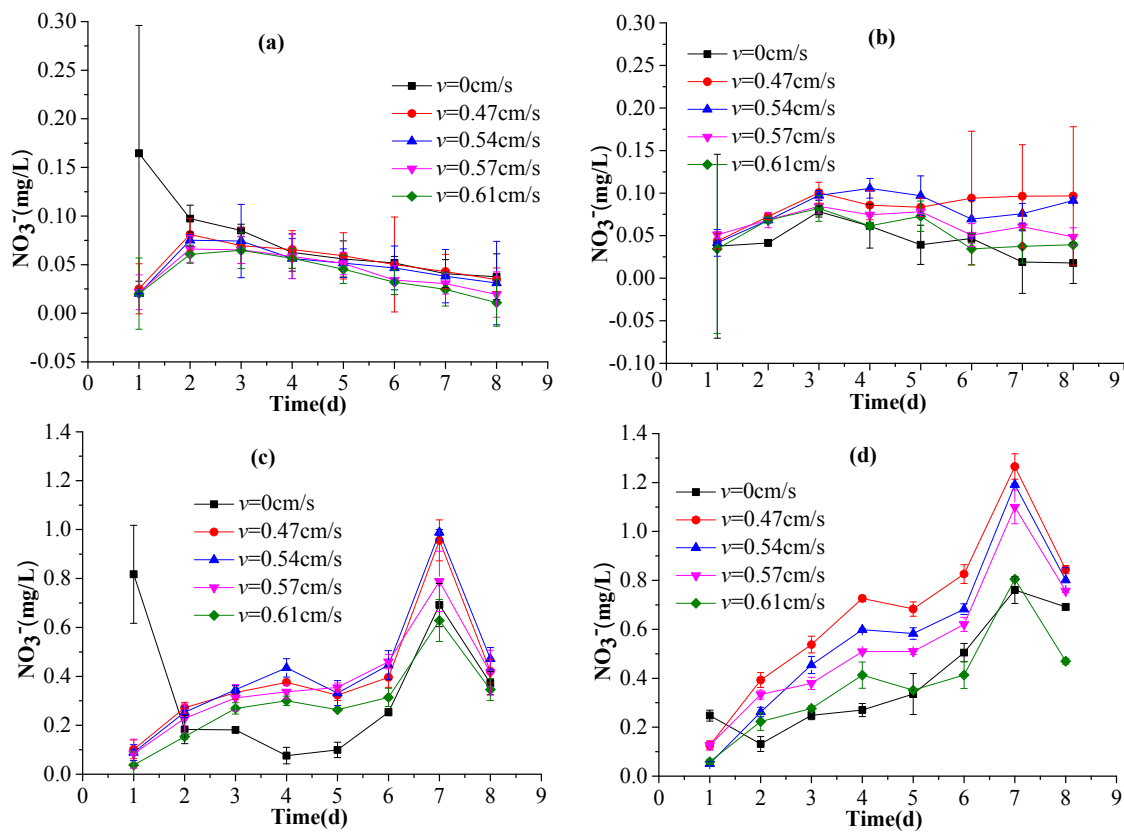


Figure 8. Nitrate (NO₃-N) concentrations at depths of: (a) 6 cm; (b) 3 cm; and (c) 1 cm within the sediments; and (d) in the overlying water, during the experiment.

Table 2. The regression equations for the NO₃-N concentrations.

Sediment Depth (cm)	Velocity <i>v</i> (cm·s ⁻¹)	Regression Equation (<i>x</i> : Day; <i>y</i> : NO ₃ -N Concentration)	Coefficient of Determination R ²
1	0.47	$y = \exp(0.237 - 2.338/x)$	0.955
	0.54	$y = \exp(0.286 - 3.263/x)$	0.973
	0.57	$y = \exp(-0.06 - 2.086/x)$	0.894
	0.61	$y = \exp(-0.32 - 2.523/x)$	0.922
Overlying Water	0.47	$y = \exp(-0.454 - 1.842/x)$	0.769
	0.54	$y = \exp(-0.325 - 2.124/x)$	0.832
	0.57	$y = \exp(-0.438 - 2.075/x)$	0.886
	0.61	$y = \exp(-0.532 - 2.742/x)$	0.927

These regression equations can be expressed as the following generic form:

$$y = \exp(a - b/x) \tag{5}$$

where *y* is the NO₃-N concentration on day *x*; and *a* and *b* are regression coefficients.

3.4. Effects of Flow Rate on SRP Concentration

Under anaerobic conditions, the degradation of organic matter was positively correlated with the SRP release flux, which in turn was negatively correlated with the DO concentration [37]. For a

given sediment depth, a higher flow rate corresponded to a lower SRP concentration, whereas for a given testing day a deeper depth had a lower SRP concentration (Figure 9). The SRP concentrations in the overlying water and in the 1-cm-deep sediments increased monotonically with time, although the concentrations in the 3-cm- and 6-cm-deep sediments decreased after their initial increase. This increasing-decreasing pattern in the SRP concentrations for the 3-cm- and 6-cm-deep sediments suggests that the P release in the sediments was likely controlled by an adsorption–desorption process. At the beginning, the overlying water had a low P content and so the sediments began to release P in an attempt to restore equilibrium. Once the P content in the overlying water increased sufficiently, the P release from the sediments achieved a dynamic balance. The maximum P release from the sediments was achieved at a relatively early stage. Figure 9d shows that under the still-water condition the sediments released P into the overlying water quite rapidly, and overall the P release was affected significantly by the hydrodynamic conditions.

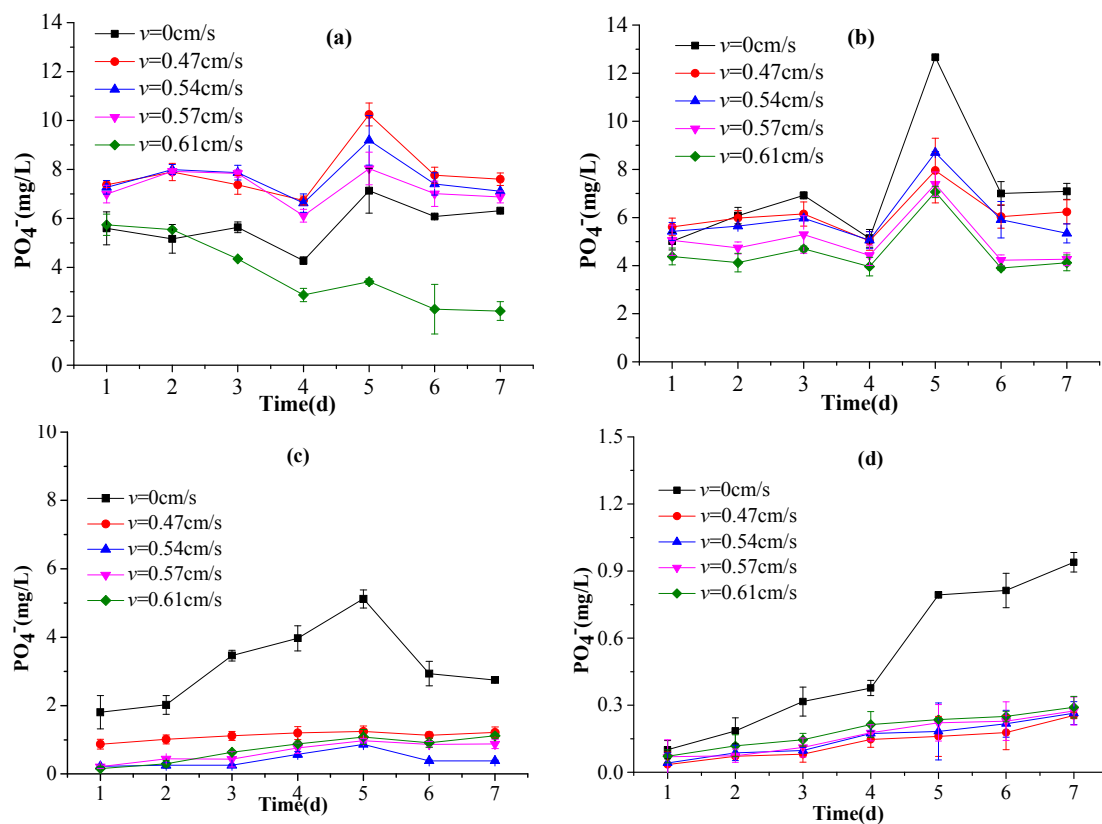


Figure 9. The concentrations of phosphate (PO_4^-) at depths of: (a) 6 cm; (b) 3 cm; and (c) 1 cm within the sediments; and (d) in the overlying water, during the experiment.

4. Discussion

The levels of $\text{NH}_3\text{-N}$, $\text{NO}_3\text{-N}$, and PO_4^- are generally used to indicate water quality in aquaculture ponds. Figure 10 shows the seven day average values of these three variables in the overlying water measured in this study under different water velocity conditions. As Figure 10a shows, the $\text{NH}_3\text{-N}$ concentrations in the overlying water decreased with increasing water velocity. When the overlying water was still, the $\text{NH}_3\text{-N}$ concentration was more than 3-fold higher than was the case when the water was moving at any of the flow rates tested, and the $\text{NH}_3\text{-N}$ concentrations decreased with increasing water velocity, although this trend slowed as the water velocity increased. Figure 10c shows that this was also true for PO_4^- : when the overlying water was still, the concentration was once again almost 3-fold higher than when the overlying water was moving, but unlike $\text{NH}_3\text{-N}$, when the water was flowing the PO_4^- concentrations increased with increasing water velocity. Similar results were reported

for earlier studies by Luo et al. [38], Duan et al. [39], Larned and Atkinson [40], and Zhu et al. [41]. For example, Larned and Atkinson found that with increasing water velocity, the rate of $\text{NH}_3\text{-N}$ and PO_4^- uptake were positively correlated with velocity by *D. cavernosa*, which indicates that the $\text{NH}_3\text{-N}$ and PO_4^- concentrations when the overlying water maintains a certain velocity are lower than those found in still overlying water. Zhu et al. found that sediment resuspension does not always increase phosphorus concentration in overlying water in shallow lakes, which means there is an effect threshold of water velocity on the lowest or highest concentrations of phosphorus. In this study, we conducted experimental measurements for a relatively coarse scale of five water velocities in the overlying water. A more detailed examination of what happens when the water velocities are within the narrow range of 0 to 0.47 cm/s, where the data shown in Figure 10 indicate major changes appear to occur, could provide useful experimental data that better explains the observed trends in nutrient concentrations.

Our findings demonstrate that in aquaculture ponds, although the water velocity is important for the removal of $\text{NH}_3\text{-N}$ and PO_4^- , it has little effect on the removal of $\text{NO}_3\text{-N}$ (Figure 10b). It does, however, provide a useful way to monitor the effects of hydrodynamic disturbance on nutrient release from the pond sediments into the overlying water.

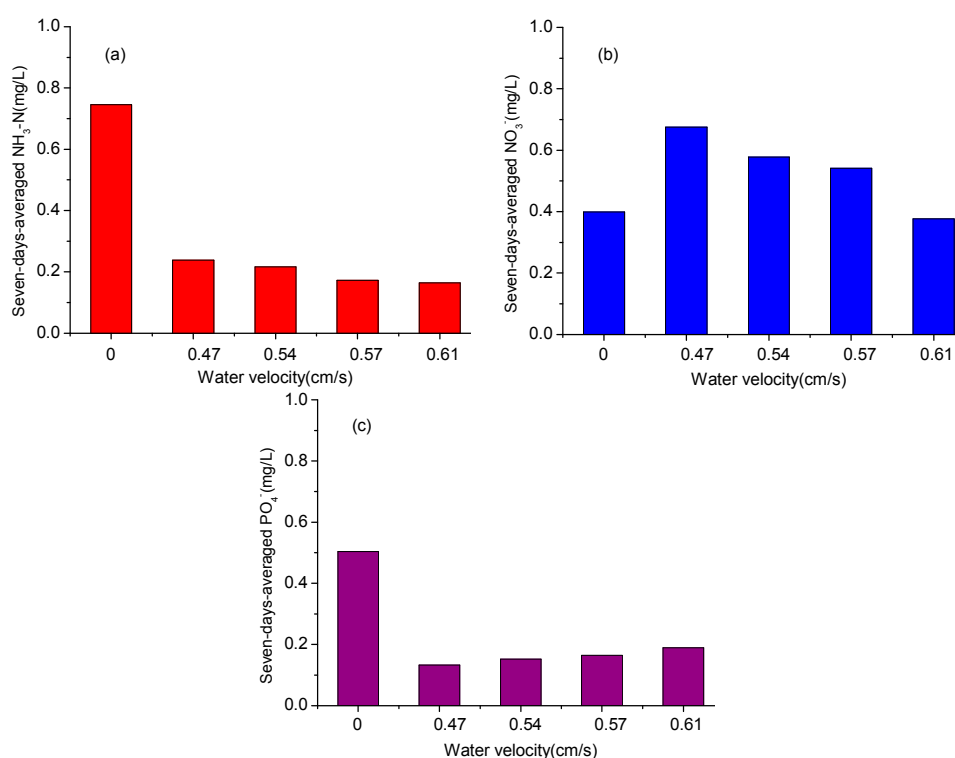


Figure 10. The seven-day-averaged concentrations in the overlying water for: (a) $\text{NH}_3\text{-N}$, (b) $\text{NO}_3\text{-N}$ and (c) PO_4^- .

5. Conclusions

This study examined the effects of flow rate on nutrient release at the aquaculture pond sediment-water interface using laboratory experiments conducted in sealed anaerobic culture model tanks. The results revealed that:

- Flow rate influenced ORP in the overlying water (pH = 6.5 to 7.0) slightly. Although the experiments were implemented under nonaerated conditions, the water could become re-aerated at certain flow rates.
- For the flowing-water conditions, a steep vertical gradient was observed in the $\text{NH}_3\text{-N}$ concentration between the overlying water and the porewater. A higher flow rate was found

to correspond to a lower $\text{NH}_3\text{-N}$ concentration in the overlying water, a larger gradient in the $\text{NH}_3\text{-N}$ concentration in the sediments, and a faster release rate of nutrients from the sediments. The trends in $\text{NH}_3\text{-N}$ concentrations were well described by a cubic function.

- The $\text{NO}_3\text{-N}$ concentrations in the overlying water and sediments tended to increase during the first five culture days, but then declined from the 6th day onward. For the still-water condition, because there was no flow turbulence $\text{NO}_3\text{-N}$ continued to be released, with some anaerobic $\text{NO}_3\text{-N}$ likely nitrified to $\text{NO}_2\text{-N}$. The trends observed in the $\text{NO}_3\text{-N}$ concentration were well described by an e-based exponential function.
- The P release from the sediments was controlled by an adsorption–desorption process. The maximum P release rate from the sediments was achieved at an early stage.

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