

6-2015

Growth in Stratospheric Chlorine from Short-Lived Chemicals not Controlled by the Montreal Protocol

R. Hossaoni

M. P. Chipperfield


A. Saiz-Lopez

J. J. Harrison

R. von Glasow

See next page for additional authors

Follow this and additional works at: https://digitalcommons.odu.edu/chemistry_fac_pubs

 Part of the [Atmospheric Sciences Commons](#), [Geochemistry Commons](#), [Geophysics and Seismology Commons](#), and the [Physical Chemistry Commons](#)

Repository Citation

Hossaoni, R.; Chipperfield, M. P.; Saiz-Lopez, A.; Harrison, J. J.; Glasow, R. von; Sommariva, R.; Atlas, E.; Navarro, M.; Montzka, S. A.; Feng, W.; and Bernath, P. F., "Growth in Stratospheric Chlorine from Short-Lived Chemicals not Controlled by the Montreal Protocol" (2015). *Chemistry & Biochemistry Faculty Publications*. 68.
https://digitalcommons.odu.edu/chemistry_fac_pubs/68

Original Publication Citation

Hossaini, R., Chipperfield, M. P., Saiz-Lopez, A., Harrison, J. J., Von Glasow, R., Sommariva, R., . . . Bernath, P. F. (2015). Growth in stratospheric chlorine from short-lived chemicals not controlled by the Montreal Protocol. *Geophysical Research Letters*, 42(11), 4573-4580. doi:10.1002/2015GL063783

Authors

R. Hossaoni, M. P. Chipperfield, A. Saiz-Lopez, J. J. Harrison, R. von Glasow, R. Sommariva, E. Atlas, M. Navarro, S. A. Montzka, W. Feng, and P. F. Bernath

RESEARCH LETTER

10.1002/2015GL063783

Key Points:

- Stratospheric Cl from short-lived chemicals has increased significantly
- Increasing Cl due to rapid growth in surface emissions of CH₂Cl₂
- COCl₂ and HCl from VSLS make a nonzero contribution to stratospheric Cl

Supporting Information:

- Texts S1 and S2, Tables S1–S9, and Figures S1–S3

Correspondence to:

R. Hossaini,
r.hossaini@see.leeds.ac.uk

Citation:

Hossaini, R., et al. (2015), Growth in stratospheric chlorine from short-lived chemicals not controlled by the Montreal Protocol, *Geophys. Res. Lett.*, 42, 4573–4580, doi:10.1002/2015GL063783.

Received 7 MAR 2015

Accepted 4 MAY 2015

Accepted article online 9 MAY 2015

Published online 1 JUN 2015

©2015. The Authors.

This is an open access article under the terms of the Creative Commons Attribution License, which permits use, distribution and reproduction in any medium, provided the original work is properly cited.

Growth in stratospheric chlorine from short-lived chemicals not controlled by the Montreal Protocol

R. Hossaini¹, M. P. Chipperfield¹, A. Saiz-Lopez², J. J. Harrison³, R. von Glasow⁴, R. Sommariva^{4,5}, E. Atlas⁶, M. Navarro⁶, S. A. Montzka⁷, W. Feng⁸, S. Dhomse¹, C. Harth⁹, J. Mühle⁹, C. Lunder¹⁰, S. O'Doherty¹¹, D. Young¹¹, S. Reimann¹², M. K. Vollmer¹², P. B. Krummel¹³, and P. F. Bernath¹⁴

¹School of Earth and Environment, University of Leeds, Leeds, UK, ²Atmospheric Chemistry and Climate Group, Institute of Physical Chemistry Rocasolano, CSIC, Madrid, Spain, ³National Centre for Earth Observation, Department of Physics and Astronomy, University of Leicester, Leicester, UK, ⁴Centre for Ocean and Atmospheric Sciences, School of Environmental Sciences, University of East Anglia, Norwich, UK, ⁵Now at Department of Chemistry, University of Leicester, Leicester, UK, ⁶Rosenstiel School of Marine and Atmospheric Science, University of Miami, Miami, Florida, USA, ⁷National Ocean and Atmospheric Administration, Boulder, Colorado, USA, ⁸National Centre for Atmospheric Science, University of Leeds, Leeds, UK, ⁹Scripps Institution of Oceanography, University of California, San Diego, California, USA, ¹⁰Monitoring and Information Technology Department, Norwegian Institute for Air Research, Kjeller, Norway, ¹¹Atmospheric Chemistry Research Group, School of Chemistry, University of Bristol, Bristol, UK, ¹²Swiss Federal Laboratories for Materials Science and Technology, Dübendorf, Switzerland, ¹³Oceans and Atmosphere Flagship, CSIRO Aspendale, Victoria, Australia, ¹⁴Department of Chemistry and Biochemistry, Old Dominion University, Norfolk, Virginia, USA

Abstract We have developed a chemical mechanism describing the tropospheric degradation of chlorine containing very short-lived substances (VSLS). The scheme was included in a global atmospheric model and used to quantify the stratospheric injection of chlorine from anthropogenic VSLS (Cl_y^{VSLS}) between 2005 and 2013. By constraining the model with surface measurements of chloroform (CHCl₃), dichloromethane (CH₂Cl₂), tetrachloroethene (C₂Cl₄), trichloroethene (C₂HCl₃), and 1,2-dichloroethane (CH₂CICH₂Cl), we infer a 2013 Cl_y^{VSLS} mixing ratio of 123 parts per trillion (ppt). Stratospheric injection of source gases dominates this supply, accounting for ~83% of the total. The remainder comes from VSLS-derived organic products, phosgene (COCl₂, 7%) and formyl chloride (CHClO, 2%), and also hydrogen chloride (HCl, 8%). Stratospheric Cl_y^{VSLS} increased by ~52% between 2005 and 2013, with a mean growth rate of 3.7 ppt Cl/yr. This increase is due to recent and ongoing growth in anthropogenic CH₂Cl₂—the most abundant chlorinated VSLS not controlled by the Montreal Protocol.

1. Introduction

In addition to long-lived source gases, such as chlorofluorocarbons and halons, halogenated very short-lived substances (VSLS)—with lifetimes in the lower troposphere < 6 months—are a source of stratospheric chlorine and bromine [e.g., Sturges et al., 2000; Mébarki et al., 2010]. VSLS enhance ozone (O₃) loss rates in the lower stratosphere [e.g., Salawitch et al., 2005], where O₃ perturbations exert a relatively large impact on climate [Riese et al., 2012; Saiz-Lopez et al., 2012; Hossaini et al., 2015]. Quantification of stratospheric VSLS loading has, therefore, been the objective of many recent field measurements and modeling studies [e.g., Hossaini et al., 2012; Tegtmeyer et al., 2013; Kreycky et al., 2013; Sala et al., 2014; Liang et al., 2014; Fernandez et al., 2014]. In particular, constraining oceanic emissions and the troposphere-to-stratosphere transport of natural brominated VSLS has been at the forefront, as these VSLS account for a significant portion (~25%) of total stratospheric bromine. Chlorinated VSLS contribute a smaller relative contribution to total stratospheric chlorine (at present a few percent) but are mostly of anthropogenic origin [e.g., Simmonds et al., 2006] and are not controlled by the Montreal Protocol.

The most abundant chlorinated VSLS are CH₂Cl₂ and CHCl₃ for which anthropogenic activity accounts for ~90% and ~25% of their tropospheric abundance, respectively [Montzka et al., 2011]. CHCl₃ is used in the manufacture of certain hydrofluorocarbons (HFCs) and is a by-product of water chlorination and bleaching processes. CH₂Cl₂ is a solvent used for paint removal, foam production, and as a feedstock for HFC production [Campbell et al., 2005]. Source gas injection (SGI) of these VSLS, together with relatively minor species

such as C_2Cl_4 and C_2HCl_3 , is estimated to provide 72 (50–95) parts per trillion (ppt) Cl to the stratosphere [Carpenter *et al.*, 2014]. This estimate, based mostly on aircraft VLS observations, is appropriate for 2012 and does not reflect the rapid growth in surface CH_2Cl_2 observed over the 2012–2013 period [Hossaini *et al.*, 2015].

In addition to SGI, it is hypothesized that VLS-derived phosgene ($COCl_2$) and hydrogen chloride (HCl), produced in the troposphere, also reach the stratosphere. Based on measured $COCl_2$ and HCl in the tropical tropopause layer (TTL), the estimated stratospheric product gas injection (PGI) of these gases contributes 25 (0–50) ppt Cl [Carpenter *et al.*, 2014]. However, observations alone cannot distinguish VLS-derived products from those derived from other sources and this is reflected in the large uncertainty range given on the above PGI estimate. This uncertainty has been acknowledged for some time in Ozone Assessment Reports [e.g., Law *et al.*, 2007; Montzka *et al.*, 2011; Carpenter *et al.*, 2014], though modeling work to resolve the issue has yet to be performed.

In this study, we developed a chemical degradation mechanism for chlorinated VLS. The scheme was included in a global model, and simulations were performed to (1) quantify stratospheric Cl_y^{VLS} , (2) assess the relative contribution of SGI versus PGI, and (3) examine the trend in stratospheric Cl_y^{VLS} over the 2005–2013 period.

2. Model and Experiments

TOMCAT is an off-line three-dimensional chemical transport model [Chipperfield, 2006]. The model has been widely used for studies of tropospheric composition, including previously the chemistry and transport of brominated VLS [e.g., Hossaini *et al.*, 2010, 2012, 2013]. TOMCAT is forced by meteorological parameters, including wind and temperature fields, taken from the European Centre for Medium-Range Weather Forecasts ERA-Interim reanalyses [Dee *et al.*, 2011]. The model contains parameterizations of mixing in the boundary layer [Holtslag and Boville, 1993] and moist convection [Tiedtke, 1989]. Large-scale vertical transport is calculated through divergence of the horizontal winds, and tracer advection follows the scheme of Prather [1986]. The model was run at a horizontal resolution of $2.8^\circ \times 2.8^\circ$ and with 31 vertical (σ - p) levels from the surface to ~ 30 km.

The version of TOMCAT used here contains a comprehensive tropospheric chemistry scheme including O_x - NO_y - HO_x - C_1 - C_3 nonmethane hydrocarbons, isoprene, and bromine chemistry [Breider *et al.*, 2010]. A gas-phase chlorine chemistry scheme has been added to the model along with five chlorinated VLS tracers; $CHCl_3$, CH_2Cl_2 , C_2Cl_4 , C_2HCl_3 , and CH_2ClCH_2Cl . Loss of these VLS, primarily through OH-initiated oxidation, yields a range of organic products, including phosgene, $COCl_2$, based on the general halocarbon degradation mechanism outlined in Ko *et al.* [2003]. A description of tracers and the model chemistry scheme is given in the supporting information.

In all simulations, background tropospheric chlorine was supplied from the breakdown of (i) ocean-emitted halocarbons (e.g., $CHBr_2Cl$, CH_2BrCl , and $CHBrCl_2$), whose emissions were specified [Ordóñez *et al.*, 2012], and (ii) the relatively long-lived source gas methyl chloride (CH_3Cl). The stratospheric-relevant source gases carbon tetrachloride (CCl_4) and methyl chloroform (CH_3CCl_3), not routinely included in this tropospheric model configuration, were also considered as they, in addition to VLS, are a $COCl_2$ source.

Three simulations covering the 2005–2013 period were performed. EXP1 (control run) contained no anthropogenic chlorinated VLS (only natural Cl sources). EXP2 was identical to EXP1 but also contained $CHCl_3$, CH_2Cl_2 , and C_2Cl_4 . A mixing ratio boundary condition, as opposed to specified emissions, was used to constrain their surface abundance in the model. This varied with latitude (five bands, $> 60^\circ N$, 30 – $60^\circ N$, 0 – $30^\circ N$, 0 – $30^\circ S$, and $> 30^\circ S$) and annually, based on surface observations from the NOAA and AGAGE monitoring networks [e.g., Prinn *et al.*, 2000; Montzka *et al.*, 2011]; see supporting information. With this boundary condition, the model was then used to simulate the tropospheric distribution and troposphere-to-stratosphere transport of VLS and their product gases. EXP3 was identical to EXP2 but also included C_2HCl_3 and CH_2ClCH_2Cl . Long-term surface observations of these VLS are sparse; therefore, they were only “switched on” in the model post 2010 and their surface abundance was scaled to give reasonable agreement with observed upper tropospheric mixing ratios (see section 3.1).

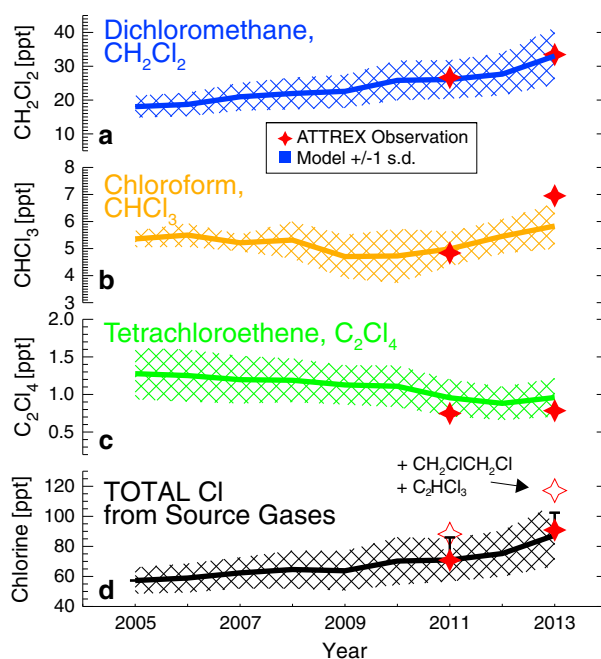


Figure 1. Time series of simulated tropical mean (a) CH_2Cl_2 , (b) CHCl_3 , and (c) C_2Cl_4 mixing ratio (ppt) at the level of zero radiative heating (LZRH) (grated area, ± 1 standard deviation) derived from observed surface values. Equivalent observed quantities at the LZRH from the 2011 (114° to 134°W longitude and 7° to $\sim 20^\circ\text{N}$ latitude, November) and 2013 (92° to 172°W longitude and 0.2° to $\sim 20^\circ\text{N}$ latitude, February/March) NASA ATTREX missions are also shown. (d) The summed total chlorine in source gases; $3 \times \text{CHCl}_3 + 2 \times \text{CH}_2\text{Cl}_2 + 4 \times \text{C}_2\text{Cl}_4$, from model run EXP2 and observed quantity shown with filled stars. The 2011 and 2013 error bars correspond to the total chlorine in source gases when $3 \times \text{C}_2\text{HCl}_3 + 2 \times \text{CH}_2\text{ClCH}_2\text{Cl}$ is also considered (i.e., EXP3, equivalent ATTREX quantity shown with open stars).

3. Results and Discussion

3.1. Source Gas Injection

Figure 1 shows the simulated time-varying abundance of CH_2Cl_2 , CHCl_3 , and C_2Cl_4 at the level of zero radiative heating (LZRH, ~ 15 km, $\sim 360\text{K}$ potential temperature) in the TTL – above which air ascends into the stratosphere [e.g., Gettelman *et al.*, 2004]. Long-term aircraft [Leedham Elvidge *et al.*, 2015] and surface observations [Hossaini *et al.*, 2015] of CH_2Cl_2 have shown a rapid increase in its tropospheric abundance in the last decade, particularly in the Northern Hemisphere (NH), owing to industrial sources [Carpenter *et al.*, 2014; Hossaini *et al.*, 2015]. This information is contained in the model forcing data and, as shown in Figure 1, caused a largely monotonic increase in the stratospheric injection of CH_2Cl_2 over the simulation period. At the LZRH, CH_2Cl_2 increased from, on average, 18 ppt in 2005 to 33 ppt in 2013, i.e., an increase of 83%.

Figure 1 also shows high-altitude CH_2Cl_2 observations from the NASA Airborne Tropical Tropopause Experiment (ATTREX). Whole air samples were collected on board the unmanned Global Hawk aircraft during deployment in 2011 and 2013 and were analyzed by the University of Miami using gas chromatography/mass spectrometry. These observations provide a valuable consistency check of the model. The agreement between modeled and observed CH_2Cl_2 is excellent. For CHCl_3 , the model exhibits a low bias of $\sim 15\%$ compared to the 2013 observations. Due to no significant long-term surface CHCl_3 trend, the stratospheric SGI of CHCl_3 varied relatively little over the simulation. C_2Cl_4 is a relatively minor contributor to chlorine SGI and its tropospheric abundance has steadily declined in the last decade [Montzka *et al.*, 2011]. Both model and observation indicate < 1 ppt of this VSLS at the LZRH in 2013.

We estimate that the stratospheric SGI of CHCl_3 , CH_2Cl_2 , and C_2Cl_4 provided a total of 87 ppt Cl in 2013. This agrees well (to within 5%) of the same quantity derived from ATTREX observations (~ 91 ppt Cl). Consideration of C_2HCl_3 and $\text{CH}_2\text{ClCH}_2\text{Cl}$ (i.e., EXP3) increases the model estimate to ~ 102 ppt Cl. The observed chlorine SGI is also greater when these gases are included; shown in Figure 1d, as the difference between the filled and open stars. This difference is virtually entirely due to $\text{CH}_2\text{ClCH}_2\text{Cl}$. Due to a particularly short local tropospheric lifetime (typically several days) and a low tropospheric abundance (< 1 ppt), C_2HCl_3 makes a negligible

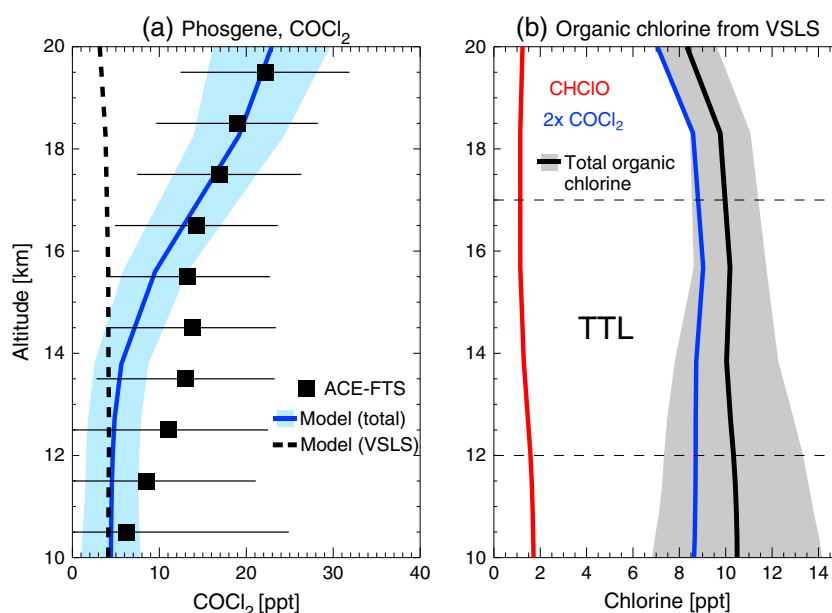


Figure 2. (a) Observed tropical mean profile of COCl_2 volume mixing ratio (ppt) from the ACE-FTS satellite instrument (2006–2012 average). The horizontal bars denote ± 1 standard deviation. Also shown are corresponding model profiles of total COCl_2 (from EXP3, solid line, shading ± 1 standard deviation) and VLS-derived COCl_2 (from EXP3 minus the control run, dashed line). (b) Simulated tropical mean profiles of organic chlorine from VLS-derived COCl_2 , CHClO , and their total (shading ± 1 standard deviation) in 2013. The location of the tropical tropopause layer (TTL) is indicated.

contribution to chlorine SGI. We find that CH_2Cl_2 , CHCl_3 , C_2Cl_4 , and $\text{CH}_2\text{ClCH}_2\text{Cl}$ account for 65%, 17%, 3.5%, and 14.5% of the total SGI of chlorine in 2013, respectively. Local lifetimes of these VLSs in the tropical boundary layer are estimated in the range 98–133, 100–136, 60–81, and 42–58 days, respectively [Carpenter *et al.*, 2014]. The contribution of $\text{CH}_2\text{ClCH}_2\text{Cl}$ could be larger as the model underestimates measured values in 2013. Long-term observations of this VLS are needed to constrain its tropospheric abundance.

Our modeled estimate of stratospheric chlorine from the SGI of VLSs in 2013 is 30% larger than that reported by Carpenter *et al.* [2014], due to the continued increases in CH_2Cl_2 mixing ratios. Note that the ocean-emitted VLSs CHBr_2Cl , CH_2BrCl , and CHBrCl_2 are not included in our estimate, as the focus of this work is on anthropogenic VLSs. Besides, their total SGI is estimated to provide < 0.4 ppt Cl [Carpenter *et al.*, 2014] and is therefore very small compared to the species considered here.

3.2. Product Gas Injection

Organic products derived from VLSs degradation include halogenated peroxy radicals, hydroperoxides, peroxy nitrates, and carbonyl compounds [e.g., Ko *et al.*, 2003; Hossaini *et al.*, 2010]. We found that the only organic products simulated with tropospheric mixing ratios > 1 ppt were (i) phosgene, COCl_2 , formed principally from CHCl_3 (with minor contributions from C_2Cl_4 and C_2HCl_3) and (ii) formyl chloride, CHClO , formed from CH_2Cl_2 . Upper tropospheric COCl_2 measurements (Figure 2) are made by the Atmospheric Chemistry Experiment Fourier transform spectrometer (ACE-FTS) on board the SCISAT satellite [e.g., Fu *et al.*, 2007; Brown *et al.*, 2011]. Here we use the v3.0 data product for measurements between 2006 and September 2010, and v3.5 for October 2010 to 2012.

Throughout the observed profile shown in Figure 2a, model-derived COCl_2 is within the measurement variability at all altitudes. In the tropical upper troposphere (~ 10 km), VLSs account for virtually all simulated COCl_2 and the agreement between model and observation is good. In the lower part of the TTL, however, model COCl_2 exhibits a low bias. Potential explanations for this include a missing source (e.g., from a VLS not considered) or an underestimate of the phosgene yield from C_2Cl_4 or C_2HCl_3 . The former seems unlikely as the most abundant chlorinated VLS known to produce phosgene, i.e., CHCl_3 , is included, though the underestimate of CHCl_3 , shown in Figure 1 (albeit small) may contribute to the low COCl_2 bias. Alternatively, a dynamical influence may include an underestimate of the in-mixing of phosgene-containing air from the extratropical lower stratosphere into the TTL.

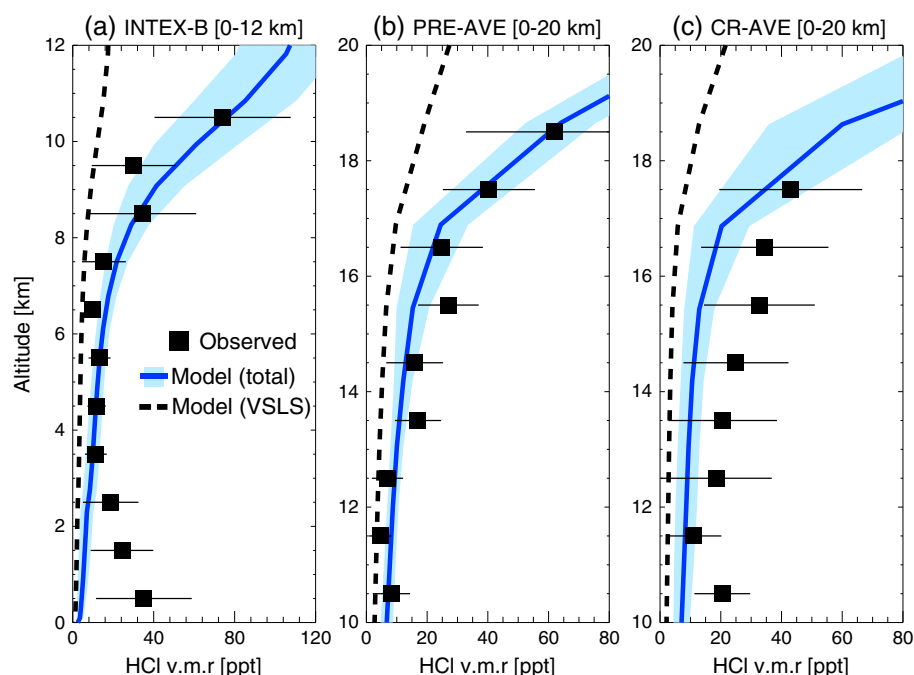


Figure 3. Observed profiles of HCl volume mixing ratio (ppt) from the NASA (a) 2006 INTEX-B mission, (b) 2004 Pre-AVE mission, and (c) 2006 CR-AVE mission. INTEX-B measurements obtained on board the DC-8 aircraft (0–12 km) between $\sim 40^\circ\text{N}$ and 61°N latitude (Anchorage deployment, Alaska). Pre-AVE and CR-AVE measurements obtained on board the WB-57 aircraft (10–20 km) within the latitude ranges $\sim 3^\circ\text{S}$ to $\sim 8^\circ\text{N}$ and $\sim 1^\circ\text{S}$ to $\sim 20^\circ\text{N}$. Measured HCl profiles are campaign means calculated in 1 km altitude bins. Horizontal lines on observed data denote ± 1 standard deviation. Also shown are corresponding model mean profiles of HCl derived from all sources (solid profile, ± 1 standard deviation) and HCl derived from anthropogenic VLS only (dashed line).

Systematic errors in the ACE-FTS COCl_2 mixing ratios are dominated by spectroscopic errors, assumed $\sim 30\%$ [Fu *et al.*, 2007]; this results from the lack of hot bands in the COCl_2 spectroscopic line list and uncertainties in line intensities. It is therefore particularly challenging to assess the fidelity of simulated COCl_2 in the TTL. We note that around the LZRH (~ 15 km), model COCl_2 agrees well with the observed values and here VLS account for $\sim 95\%$ of total COCl_2 . Above 20 km, COCl_2 production is dominated by the long-lived solvents CCl_4 and CH_3CCl_3 .

To our knowledge, no atmospheric observations of the second most abundant product gas, CHClO , exist. Its simulated surface mixing ratio in the NH, where its primary source in the model (CH_2Cl_2) is most abundant, is ~ 3 – 5 ppt (see supporting information). Figure 2b shows that CHClO provides a small, though previously unidentified, contribution to PGI of ~ 2 ppt Cl.

In addition to organic products, VLS-derived HCl may also enter the stratosphere through PGI. In situ measurements of tropospheric HCl are generally sparse, though measured profiles are available from previous NASA aircraft missions [Marcy *et al.*, 2004, 2007; Kim *et al.*, 2008], shown in Figure 3, and more recent aircraft sampling around the tropopause [Jurkat *et al.*, 2014]. In the marine boundary layer, sea salt is the dominant chlorine source [e.g., Saiz-Lopez and von Glasow, 2012]. Here simulated HCl exhibits a low bias against observed values as sea salt chlorine is not considered (as not relevant for the stratosphere). However, in the free troposphere and TTL—where chlorine is mostly derived from the breakdown of organic source gases—agreement between simulated and observed HCl is generally good. In 2013, VLS accounted for, on average, $\sim 40\%$ of total HCl around the LZRH (~ 9 ppt of VLS-derived HCl in absolute terms), with the remainder supplied from a combination of stratospheric-influenced air and the in situ oxidation of the relatively long-lived source gas CH_3Cl [e.g., Marcy *et al.*, 2007].

By summing the modeled organic and inorganic contributions, we estimate a total stratospheric PGI of ~ 21 ppt Cl in 2013 (EXP3). This value falls within the estimated range of 0–50 ppt Cl reported by Carpenter *et al.* [2014] but is slightly lower than their best estimate of 25 ppt Cl. This difference is largely due to the lower

Table 1. Simulated Source and Product Gas Contributions to Total Stratospheric Chlorine Supplied From VSLS (ppt Cl)^a

	2005	2006	2007	2008	2009	2010	2011	2012	2013	WMO 2014 ^c
Source Gas Sum ^b	57.3	59.0	62.4	64.6	63.8	70.3	70.9	75.2	87.4 [102.3]	72(50–95)
Phosgene	8.0	7.7	8.1	8.8	7.9	9.0	8.6	8.9	8.3 [9.1]	15(0–30)
Formyl Chloride	1.1	1.1	1.3	1.3	1.4	1.6	1.5	1.6	2.4 [2.4]	Not considered
Hydrogen Chloride	3.0	2.9	3.2	3.6	3.0	3.3	3.3	3.1	7.6 [9.4]	10(0–20)
Product Gas Sum	12.1	11.8	12.7	13.8	12.3	13.9	13.5	13.7	18.4 [20.9]	25(0–50)
TOTAL Cl	69.4	70.8	75.1	78.4	76.1	82.4	84.4	88.9	105.8 [123.2]	95(50–145)

^aModel estimates are reported from EXP2. For 2013, values are also reported for EXP3 (square brackets). Estimates from the 2014 WMO/UNEP Scientific Assessment of Ozone Depletion are compared to the simulated values from this work.

^bSum of ($3 \times \text{CHCl}_3$) + ($2 \times \text{CH}_2\text{Cl}_2$) + ($4 \times \text{C}_2\text{Cl}_4$); with additional contribution of ($3 \times \text{C}_2\text{HCl}_3$) and ($2 \times \text{CH}_2\text{ClCH}_2\text{Cl}$) (see data in square brackets).

^cBest estimate and range reported in WMO/UNEP Scientific Assessment of Ozone Depletion 2014 [Carpenter *et al.*, 2014].

PGI contribution from COCl_2 from the model relative to the estimate of Carpenter *et al.* [2014], as summarized in Table 1. Note that our total PGI estimate also includes a contribution from CHClO which has been previously unquantified. Our analysis shows, for the first time, that PGI due to VSLS-derived products constitutes a nonzero supply of chlorine to the stratosphere. However, further insight into the impact of mechanistic and kinetic uncertainties in the chemical mechanism employed is required, beyond the scope of this work, in order to fully constrain the derived chlorine budget. In particular, this includes a need for knowledge of the aqueous phase processing of soluble product gases (e.g., COCl_2 , CHClO) and heterogeneous processes occurring on, for example, cirrus ice particles within the TTL [e.g., von Hobe *et al.*, 2011].

3.3. Total Chlorine

We estimate the total chlorine injection from anthropogenic VSLS ($\text{Cl}_y^{\text{VSLS}}$)—the sum of SGI and PGI contributions—was 123 ppt Cl in 2013 (EXP3, Table 1). SGI accounts for $\sim 83\%$ of total $\text{Cl}_y^{\text{VSLS}}$ with the remainder supplied from PGI of VSLS-derived COCl_2 (7%), CHClO (2%), and HCl (8%). Our simulated $\text{Cl}_y^{\text{VSLS}}$ is larger than the best estimate of 95 ppt Cl reported by Carpenter *et al.* [2014], owing to a larger SGI contribution and the continued increase in CH_2Cl_2 mixing ratios. However, it falls within their reported 50–145 ppt Cl range.

To estimate the purely anthropogenic component of stratospheric $\text{Cl}_y^{\text{VSLS}}$, a further sensitivity experiment (EXP4) was performed, identical to EXP3, but with surface CH_2Cl_2 and CHCl_3 reduced by 10% and 75%, respectively (to remove the estimated natural component of emissions). C_2Cl_4 , C_2HCl_3 , and $\text{CH}_2\text{ClCH}_2\text{Cl}$ are thought to be exclusively anthropogenic and were therefore unchanged. From EXP4 we estimate $\sim 80\%$ of the simulated stratospheric $\text{Cl}_y^{\text{VSLS}}$ in 2013 is attributable to anthropogenic activity.

Based on EXP2, as it contained the same VSLS throughout the simulation, $\text{Cl}_y^{\text{VSLS}}$ increased by $\sim 52\%$ over the 2005–2013 period (Table 1). A linear fit to the data indicates a mean growth rate of ~ 3.7 ppt Cl yr^{-1} (Figure 4),

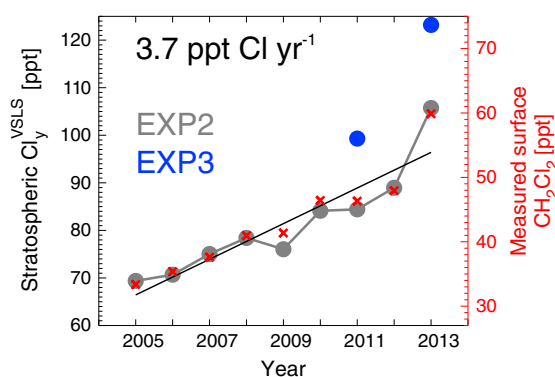


Figure 4. Time series of simulated annual mean total stratospheric chlorine from VSLS (sum of source and product gas contributions, filled circles, left axis). A linear trend line is applied to EXP2 which considered CH_2Cl_2 , CHCl_3 , and C_2Cl_4 , and the 2005–2013 mean growth rate (ppt Cl^{-1}) is annotated. EXP3 (blue) also considered C_2HCl_3 and $\text{CH}_2\text{ClCH}_2\text{Cl}$. Also overlaid are annual mean CH_2Cl_2 mixing ratios (ppt) in the Northern Hemisphere calculated from NOAA surface observations (red crosses, right axis).

driven by the ongoing increase in the surface concentration of CH_2Cl_2 . In addition to causing an enhanced SGI, CH_2Cl_2 growth has also enhanced PGI; both CHClO (derived from CH_2Cl_2) and HCl exhibited a sharp increase over the 2012–2013 period, coinciding with a particularly sharp increase in surface CH_2Cl_2 and indeed total $\text{Cl}_y^{\text{VLSL}}$ (Figure 4). Between 2012 and 2013, $\text{Cl}_y^{\text{VLSL}}$ increased by $\sim 20\%$.

4. Concluding Remarks

Knowledge of the stratospheric loading of ozone-depleting substances and their trends is required to predict the future evolution of stratospheric O_3 and recovery of the ozone layer. Constraining the supply of chlorine from chemicals not regulated by the Montreal Protocol is particularly important. Here, using a global model supported by atmospheric observations, we show that the contribution of anthropogenic VLSL has increased significantly in the last decade. A 2013 stratospheric $\text{Cl}_y^{\text{VLSL}}$ loading of > 100 ppt is large enough to significantly impact O_3 loss rates in the lower stratosphere [Hossaini et al., 2015], where surface temperature and climate are particularly sensitive to O_3 perturbations. Indeed, trends in anthropogenic chlorine VLSL are suggested to already have contributed a nonzero amount to the radiative forcing of climate since the preindustrial era [Hossaini et al., 2015].

Although $\text{Cl}_y^{\text{VLSL}}$ remains small in comparison to the total stratospheric chlorine load from long-lived ozone-depleting substances (> 3000 ppt Cl), it is far larger than that supplied from newly detected chlorofluorocarbons and hydrochlorofluorocarbons (< 10 ppt Cl) [Laube et al., 2014]. Further, $\text{Cl}_y^{\text{VLSL}}$ will continue to increase if the observed growth in tropospheric CH_2Cl_2 continues. This could likely be the case if the upward CH_2Cl_2 trend is attributable to the continued industrialization of developing countries and/or release following its use as a feedstock in the production of HFCs, the second generation of CFC replacement gases, for which a global market expansion seems likely in the coming years, or indeed other chemicals.

Acknowledgments

We thank the Natural Environment Research Council (NERC) for funding through the TropHAL project (NE/J02449X/1 and NE/J022780/1). We thank Ron Prinn, Ray Weiss, and their AGAGE colleagues for data access. AGAGE is supported principally by NASA (USA) grants to MIT and SIO and also by DECC (UK) and NOAA (USA) grants to Bristol University; CSIRO and BoM (Australia); Empa (Switzerland); and NILU (Norway). S.A.M. acknowledges the technical assistance of C. Siso, B. Miller, and B. Hall and support, in part, from NOAA Climate Program Office's AC4 Program. The ACE mission is supported by the Canadian Space Agency. R.H. thanks D. Fahey and R. Gao for useful discussion. Model output data presented here are available on request. Surface observation data are given in the supporting information.

The Editor thanks two anonymous reviewers for their assistance in evaluating this paper.

References

- Breider, T. J., M. P. Chipperfield, N. A. D. Richards, K. S. Carslaw, G. W. Mann, and D. V. Spracklen (2010), Impact of BrO on dimethylsulfide in the remote marine boundary layer, *Geophys. Res. Lett.*, *37*, L02807, doi:10.1029/2009GL040868.
- Brown, A. T., M. P. Chipperfield, C. Boone, C. Wilson, K. A. Walker, and P. F. Bernath (2011), Trends in atmospheric halogen containing gases since 2004, *J. Quant. Spectrosc. Radiat. Transfer*, *112*(16), 2552–2566, doi:10.1016/j.jqsrt.2011.07.005.
- Campbell, N., et al. (2005), HFCs and PFCs: Current and future supply, demand and emissions, plus emissions of CFCs, HCFCs and halons, in *Safeguarding the Ozone Layer and the Global Climate System, Issues Related to Hydrofluorocarbons and Perfluorocarbons, IPCC/TEAP Special Report*, chap. 11, pp. 404–436, Cambridge, Univ. Press, Cambridge, U. K.
- Carpenter, L., S. Reimann, J. B. Burkholder, C. Clerboux, B. D. Hall, R. Hossaini, J. C. Laube, and S. A. Yvon-Lewis (2014), Ozone-depleting substances (ODSs) and other gases of interest to the Montreal Protocol, in *Scientific Assessment of Ozone Depletion: 2014, Global Ozone Research and Monitoring Project - Report No. 55*, edited by A. Engel and S. A. Montzka, chap. 1, pp. 1–101, World Meteorol. Organ., Geneva, Switzerland.
- Chipperfield, M. P. (2006), New version of the TOMCAT/SLIMCAT off-line chemical transport model: Intercomparison of stratospheric tracer experiments, *Q. J. R. Meteorol. Soc.*, *132*, 1179–1203, doi:10.1256/qj.05.51.
- Dee, D. P., et al. (2011), The ERA-Interim reanalysis: Configuration and performance of the data assimilation system, *Q. J. R. Meteorol. Soc.*, *137*, 553–597, doi:10.1002/qj.828.
- Fernandez, R. P., R. J. Salawitch, D. E. Kinnison, J.-F. Lamarque, and A. Saiz-Lopez (2014), Bromine partitioning in the tropical tropopause layer: Implications for stratospheric injection, *Atmos. Chem. Phys.*, *14*, 13,391–13,410, doi:10.5194/acp-14-13391-2014.
- Fu, D., C. D. Boone, P. F. Bernath, K. A. Walker, R. Nassar, G. L. Manney, and S. D. McLeod (2007), Global phosgene observations from the atmospheric chemistry experiment (ACE) mission, *Geophys. Res. Lett.*, *34*, L17815, doi:10.1029/2007GL029942.
- Gettelman, A., P. M. d. F. Forster, M. Fujiwara, Q. Fu, H. Vömel, L. K. Gohar, C. Johanson, and M. Ammerman (2004), Radiation balance of the tropical tropopause layer, *J. Geophys. Res.*, *109*, D07103, doi:10.1029/2003JD004190.
- Holtslag, A., and B. Boville (1993), Local versus nonlocal boundary-layer diffusion in a global climate model, *J. Clim.*, *6*, 1825–1842.
- Hossaini, R., et al. (2013), Evaluating global emission inventories of biogenic bromocarbons, *Atmos. Chem. Phys.*, *13*, 11,819–11,838, doi:10.5194/acp-13-11819-2013.
- Hossaini, R., M. P. Chipperfield, B. M. Monge-Sanz, N. A. D. Richards, E. Atlas, and D. R. Blake (2010), Bromoform and dibromomethane in the tropics: A 3-D model study of chemistry and transport, *Atmos. Chem. Phys.*, *10*, 719–735, doi:10.5194/acp-10-719-2010.
- Hossaini, R., M. P. Chipperfield, W. Feng, T. J. Breider, E. Atlas, S. A. Montzka, B. R. Miller, F. Moore, and J. Elkins (2012), The contribution of natural and anthropogenic very short-lived species to stratospheric bromine, *Atmos. Chem. Phys.*, *12*, 371–380, doi:10.5194/acp-12-371-2012.
- Hossaini, R., M. P. Chipperfield, S. A. Montzka, A. Rap, S. Dhomse, and W. Feng (2015), Efficiency of short-lived halogens at influencing climate through depletion of stratospheric ozone, *Nat. Geosci.*, *8*, 186–190, doi:10.1038/ngeo2363.
- Jurkat, T., et al. (2014), A quantitative analysis of stratospheric HCl, HNO_3 , and O_3 in the tropopause region near the subtropical jet, *Geophys. Res. Lett.*, *41*, 3315–3321, doi:10.1002/2013GL059159.
- Kim, S., et al. (2008), Airborne measurements of HCl from the marine boundary layer to the lower stratosphere over the North Pacific Ocean during INTEX-B, *Atmos. Chem. Phys. Discuss.*, *8*, 3563–3595, doi:10.5194/acpd-8-3563-2008.
- Ko, K., et al. (2003), Halogenated very short-lived substances, in *Scientific Assessment of Ozone Depletion: 2002, Global Ozone Res. Monit. Project-Rep. 47*, chap. 2, pp. 1–57, World Meteorol. Organ., Geneva, Switzerland.

- Kreycy, S., C. Camy-Peyret, M. P. Chipperfield, M. Dorf, W. Feng, R. Hossaini, L. Kritten, B. Werner, and K. Pfeilsticker (2013), Atmospheric test of the $J(\text{BrONO}_2)/\text{kBrO}+\text{NO}_2$ ratio: implications for total stratospheric Br_y and bromine-mediated ozone loss, *Atmos. Chem. Phys.*, *13*, 6263–6274, doi:10.5194/acp-13-6263-2013.
- Laube, J. C., et al. (2014), Newly detected ozone-depleting substances in the atmosphere, *Nat. Geosci.*, *7*, 266–269, doi:10.1038/ngeo2109.
- Law, K., et al. (2007), Halogenated very short-lived substances, in *Scientific Assessment of Ozone Depletion: 2006, Global Ozone Res. Monit. Project-Rep. 50*, chap. 2, 572 pp., World Meteorol. Organ., Geneva, Switzerland.
- Leedham Elvidge, E. C., D. E. Oram, J. C. Laube, A. K. Baker, S. A. Montzka, S. Humphrey, D. A. O'Sullivan, and C. A. M. Brenninkmeijer (2015), Increasing concentrations of dichloromethane, CH_2Cl_2 , inferred from CARIBIC air samples collected 1998–2012, *Atmos. Chem. Phys.*, *15*, 1939–1958, doi:10.5194/acp-15-1939-2015, 2015.
- Liang, Q., E. Atlas, D. Blake, M. Dorf, K. Pfeilsticker, and S. Schauffler (2014), Convective transport of very short lived bromocarbons to the stratosphere, *Atmos. Chem. Phys.*, *14*, 5781–5792, doi:10.5194/acp-14-5781-2014.
- Marcy, T., et al. (2007), Measurements of trace gases in the tropical tropopause layer, *Atmos. Env.*, *41*, 7253–7261, doi:10.1016/j.atmosenv.2007.05.032.
- Marcy, T. P., et al. (2004), Quantifying stratospheric ozone in the upper troposphere with in situ measurements of HCl, *Science*, *304*, 261–265, doi:10.1126/science.1093418.
- Mébarki, Y., V. Catoire, N. Huret, G. Berthet, C. Robert, and G. Poulet (2010), More evidence for very short-lived substance contribution to stratospheric chlorine inferred from HCl balloon-borne in situ measurements in the tropics, *Atmos. Chem. Phys.*, *10*, 397–409, doi:10.5194/acp-10-397-2010.
- Montzka, S., et al. (2011), Ozone-depleting substances (ODSs) and related chemicals, in *Scientific Assessment of Ozone Depletion: 2010, Global Ozone Res. and Monit. Proj.-Rep. 52.*, chap. 1, pp. 1–112, World Meteorol. Organ., Geneva, Switzerland.
- Montzka, S. A., M. Krol, E. Dlugokencky, B. Hall, P. Jöckel, and J. Lelieveld (2011), Small interannual variability of global atmospheric hydroxyl, *Science*, *331*, 67–69, doi:10.1126/science.1197640.
- Ordóñez, C., J.-F. Lamarque, S. Tilmes, D. E. Kinnison, E. L. Atlas, D. R. Blake, G. S. Santos, G. Brasseur, and A. Saiz-Lopez (2012), Bromine and iodine chemistry in a global chemistry-climate model: Description and evaluation of very short-lived oceanic sources, *Atmos. Chem. Phys.*, *12*, 1423–1447, doi:10.5194/acp-12-1423-2012.
- Prather, M. (1986), Numerical advection by conservation of second-order moments, *J. Geophys. Res.*, *91*(D6), 6671–6681, doi:10.1029/JD091iD06p06671.
- Prinn, R. G., et al. (2000), A history of chemically and radiatively important gases in air deduced from ALE/GAGE/AGAGE, *J. Geophys. Res.*, *105*(D14), 17,751–17,792, doi:10.1029/2000JD900141.
- Riese, M., F. Ploeger, A. Rap, B. Vogel, P. Konopka, M. Dameris, and P. Forster (2012), Impact of uncertainties in atmospheric mixing on simulated UTLS composition and related radiative effects, *J. Geophys. Res.*, *117*, D16305, doi:10.1029/2012JD017751.
- Saiz-Lopez, A., and R. von Glasow (2012), Reactive halogen chemistry in the troposphere, *Chem. Soc. Rev.*, *41*, 6448–6472, doi:10.1039/c2cs35208g.
- Saiz-Lopez, A., et al. (2012), Estimating the climate significance of halogen-driven ozone loss in the tropical marine troposphere, *Atmos. Chem. Phys.*, *12*, 3939–3949, doi:10.5194/acp-12-3939-2012.
- Sala, S., H. Bönisch, T. Keber, D. E. Oram, G. Mills, and A. Engel (2014), Deriving an atmospheric budget of total organic bromine using airborne in situ measurements from the western Pacific area during SHIVA, *Atmos. Chem. Phys.*, *14*, 6903–6923, doi:10.5194/acp-14-6903-2014.
- Salawitch, R., D. Weisenstein, L. Kovalenko, C. Sioris, P. Wennberg, K. Chance, M. Ko, and C. McLinden (2005), Sensitivity of ozone to bromine in the lower stratosphere, *Geophys. Res. Lett.*, *32*, L05811, doi:10.1029/2004GL021504.
- Simmonds, P. G., et al. (2006), Global trends, seasonal cycles, and European emissions of dichloromethane, trichloroethene, and tetrachloroethene from the AGAGE observations at Mace Head, Ireland, and Cape Grim, Tasmania, *J. Geophys. Res.*, *111*, D18304, doi:10.1029/2006JD007082.
- Sturges, W., D. Oram, L. Carpenter, S. Penkett, and A. Engel (2000), Bromoform as a source of stratospheric bromine, *Geophys. Res. Lett.*, *27*(14), 2081–2084, doi:10.1029/2000GL011444.
- Tegtmeier, S., et al. (2013), The contribution of oceanic methyl iodide to stratospheric iodine, *Atmos. Chem. Phys.*, *13*, 11,869–11,886, doi:10.5194/acp-13-11869-2013.
- Tiedtke, M. (1989), A comprehensive mass flux scheme for cumulus parameterization in large-scale models, *Mon. Weather Rev.*, *117*(8), 1779–1800, doi:10.1175/1520-0493.
- von Hobe, M., et al. (2011), Evidence for heterogeneous chlorine activation in the tropical UTLS, *Atmos. Chem. Phys.*, *11*, 24–256, doi:10.5194/acp-11-241-2011.