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# Changes in the sea-ice brine community during the spring-summer transition, McMurdo Sound, Antarctica. II. Phagotrophic protists

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**ABSTRACT:** The land-fast sea-ice brine contains a diverse phagotrophic protist assemblage consisting of <math>5\ \mu\text{m}</math> heterotrophic flagellates, *Cryothecomonas* spp., heterotrophic dinoflagellates, and heterotrophic and mixotrophic ciliates. Fine-scale horizontal spatial variability is a feature of this assemblage; samples taken within 1 m of each other can be dominated by different heterotrophic protists. Many of the larger heterotrophic protists found in the brine are also found in the water column. The photosynthetic ciliate *Mesodinium rubrum* is also common. In mid to late austral spring, the heterotrophic assemblage accounts for ca 10% of the total protist biomass in the brine and is dominated by *Cryothecomonas* spp. This flagellate can reach densities of over  $10^6$  cells  $\text{l}^{-1}$  of brine. In the early austral summer, ciliates (primarily *Strombidium* spp., *Mesodinium rubrum* and *Didinium* spp.) and heterotrophic dinoflagellates (primarily a small *Gymnodinium* sp. and *Polykrikos* sp.) increase in abundance in the brine. Ciliate densities of  $\geq 3 \times 10^3$   $\text{l}^{-1}$  and heterotrophic dinoflagellate densities of  $10^4$  cells  $\text{l}^{-1}$  are common in the brine during early summer. By the end of January (just prior to ice decay and break-out), heterotrophic flagellates and ciliates can account for 50% of the protist biomass.

## INTRODUCTION

Sea-ice is a dominant feature of polar seas; it annually covers a maximum of ca  $20 \times 10^6$   $\text{km}^2$  in the southern hemisphere (Zwally et al. 1983) and ca  $14 \times 10^6$   $\text{km}^2$  in the northern hemisphere (Comiso 1986). Sea-ice plays an important role in determining energy balance and ocean-atmosphere interactions on a global scale (reviewed in Eicken 1992, Legendre et al. 1992). It is a major habitat for marine biota in polar environments (Garrison et al. 1986, Horner et al. 1992) and is important in structuring polar marine ecosystems (Eicken 1992).

The sea-ice microbial community is important in several respects. Production by sea-ice protists can make a significant contribution to food webs in the plankton and in the benthos (Knox 1990, Matsuda et al. 1990, Bianchi et al. 1992, Lizotte & Sullivan 1992). Furthermore, the presence and growth of protists in surface and interior habitats in sea-ice is thought to: (1) increase light attenuation in the sea-ice and alter

the spectral quality of light reaching the base of the sea-ice and the underlying water column (SooHoo et al. 1987, Arrigo et al. 1991), and (2) increase heat absorption and thus influence the physical structure of sea-ice and its break-up (Meguro 1962, Buynitskiy 1968, McConville & Wetherbee 1983, Maykut 1985, Eicken et al. 1991, Stoecker et al. 1992).

In polar environments 2 types of sea-ice occur. Land-fast sea-ice forms and remains fast along the coast, the rest is pack-ice (Horner et al. 1992). In general, the 2 types differ in their mode of formation, physical structure and biota (reviewed in Garrison et al. 1986, Horner et al. 1992). In land-fast ice, protist communities are often most conspicuous at or near the bottom of the sea-ice and are diatom-dominated (Hoshiai 1972, McConville & Wetherbee 1983, Palmisano & Sullivan 1983, Sasaki & Watanabe 1984). In pack-ice, sub-surface brine and interior algal assemblages are most common (Ackley et al. 1979, Garrison & Buck 1989, 1991, Spindler et al. 1990). Pack-ice communities are often dominated by diatoms, but also contain a diverse

assemblage of other autotrophs and phagotrophic protists (Garrison & Buck 1989, Spindler et al. 1990).

Recently we reported a flagellate-dominated brine assemblage from the upper land-fast ice in McMurdo Sound, Antarctica (Stoecker et al. 1990, 1991, 1992). During late December and early January, the autotrophic assemblage in the brine peaks, with biomass reaching  $\geq 100 \mu\text{g l}^{-1}$  and chlorophyll *a*  $\geq 3 \mu\text{g l}^{-1}$  (Stoecker et al. 1992). This assemblage is distinct from the diatom-dominated assemblages at the base of the sea-ice. The dominant photosynthetic protists in the upper land-fast brine are a small non-thecate dinoflagellate, a prasinomonad (*Mantoniella* sp.), and unidentified chrysophyte statocysts (Stoecker et al. 1992); these taxa are also common in pack-ice (Garrison & Buck 1989).

Although heterotrophic protists are often the major consumers of primary production and the primary agents of nutrient regeneration in aquatic environments (reviewed in Capriulo et al. 1991, Caron 1991), almost nothing was known about their ecology in land-fast sea-ice. Algal and bacterial assemblages have been studied in land-fast sea-ice (Kottmeier & Sullivan 1988), but the heterotrophic protists in this habitat have largely been ignored. Preliminary studies in 1989–90 revealed that ciliates were present in the upper land-fast sea-ice brine community in McMurdo Sound (Stoecker et al. 1990).

In pack-ice surface brine assemblages, heterotrophic protists can comprise from 1 to >93 % of the protist biomass; it has been hypothesized that grazing by protists influences algal growth and biomass accumulation (Garrison & Buck 1991). Studies in pack-ice indicate that an active microbial food web functions within sea-ice, but time series data on the population dynamics of autotrophic and heterotrophic components of the community were not available (Garrison & Buck 1989). The occurrence of very high ammonia concentrations in some samples of both land-fast and pack-ice suggest that the activities of heterotrophic protists may have an important influence on biological and chemical parameters in sea-ice (Arrigo et al. 1990, Dieckmann et al. 1991).

Here we report on the heterotrophic and mixotrophic components of the protist assemblage in the upper land-fast in McMurdo Sound during the spring-summer transition. Access to the land-fast ice at McMurdo allows repeated observation of the upper sea-ice brine community, and thus seasonal trends can be investigated. Because the upper land-fast ice brine assemblage has many properties in common with the more complex and more difficult to sample pack-ice assemblage, it may be an excellent model system in which to observe processes common to both land-fast and pack-ice communities (Stoecker et al. 1992).

## METHODS

Samples were collected by drilling holes approximately 50 cm into the land-fast, annual sea-ice in McMurdo Sound, Antarctica (Fig. 1). Brine was allowed to accumulate in the holes, and the accumulated brine was collected by gently pumping it with a hand vacuum pump into a polycarbonate bottle wrapped with duct tape to keep the contents from being exposed to light (Stoecker et al. 1992). During the 1990–91 field season, 11 sites 25 to 50 m from the retreating ice edge were sampled between 15 December and 31 January. Sampling at the ice edge sites was biased by the fact that we only sampled sites where the sea-ice was suitable for a helicopter to land. Since the ice edge was retreating and thus constantly changing in position, each site was only sampled once. During our field season, the ice edge on the eastern side of the sound retreated from north of Cape Byrd to Cape Royds; the area in which we sampled this retreating edge is shown in Fig. 1. For most sites/dates, between 2 and 4 replicate sampling holes were drilled at 1 m intervals along a short transect on the ice.

For comparison with the ice edge sites and to allow for repeated sampling at 1 site, a fixed location (H90) on the annual land-fast sea-ice near McMurdo Station was chosen for sampling (Fig. 1). Contrary to our expectations, the sea-ice at H90 decayed faster than at some more northern locations in McMurdo Sound, and after early January 1991, it was not possible to access this site. Triplicate brine samples, spaced at 1 m intervals, were taken on 3 dates between 30 November and 5 January at H90.

At both the ice edge and H90 sites, brine temperature was determined by immersing a thermometer in the brine which had accumulated in the holes. Brine samples stored in insulated containers were taken to the laboratory within a few hours of collection where samples were fixed for later microscopic enumeration. Salinity was determined in the laboratory with a refractometer. Triplicate subsamples were extracted in 90 % acetone and the chlorophyll *a* content determined by fluorometry (Parsons et al. 1984).

For enumeration and sizing of <20  $\mu\text{m}$  protists, 5 to 20 ml subsamples were preserved with glutaraldehyde (final conc. 0.3 %), stained with proflavine (final conc.  $0.5 \mu\text{g ml}^{-1}$ ) and collected on 1.0  $\mu\text{m}$  pore size polycarbonate filters and examined using epifluorescent microscopy (BP 450 to 490 nm excitation filter, FT 510 chromatic beam splitter, Lp 520 nm barrier filter). With this technique, cells can be categorized as plastidic or non-plastidic based on their fluorescence patterns (Haas 1982). Protists >20  $\mu\text{m}$  in size were preserved either in 5 % acid Lugol's solution or 5 % (final conc.) buffered formalin, 50 or 100 ml samples were settled

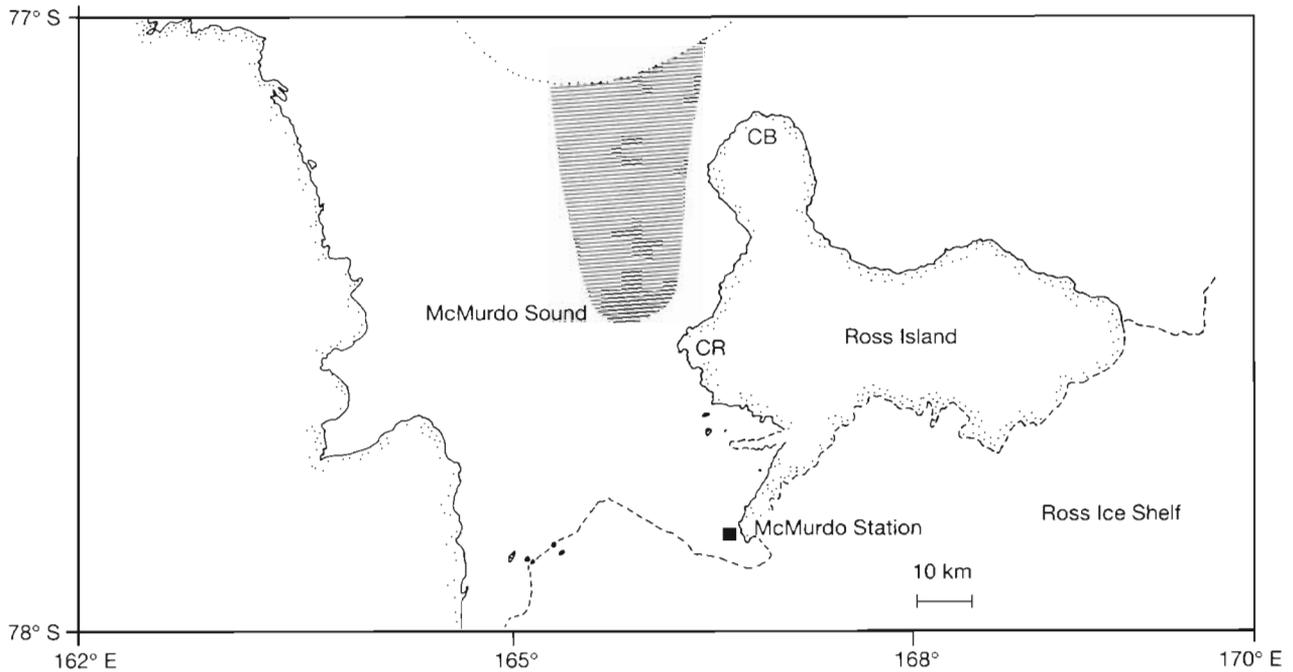


Fig. 1. Sampling locations on the annual land-fast sea-ice in McMurdo Sound, Antarctica, 1990–91. During the sampling season, the ice edge was retreating and hence constantly changing. (.....) Approximate location of the ice edge at the beginning of our sampling season in early December; (-----) the margin between the annual land-fast sea-ice and the Ross Ice Shelf. Shading indicates location of ice edge sites on the east side of McMurdo Sound. During our sampling season, the ice edge within this area retreated from the north of Cape Byrd (CB) to the vicinity of Cape Royds (CR). Ice edge land-fast ice sampling sites were approximately 25 to 50 m to the south of the retreating edge (precise locations are not known). (■) H90, located near McMurdo Station on the land-fast ice. At the times this site was sampled, the ice edge was north of Cape Royds

and examined using transmitted and epifluorescence microscopy (Stoecker et al. 1992). Sizing of  $>20 \mu\text{m}$  protists was based on specimens preserved with formalin. Cell sizes were measured with an ocular micrometer and converted to cell volumes using appropriate geometric formulae. The same conversion factor used for the microalgae (Stoecker et al. 1992),  $0.14 \text{ pg C } \mu\text{m}^{-3}$ , was used to convert protozoan biovolumes to biomass. This factor has been suggested for heterotrophic dinoflagellates fixed with glutaldehyde (Lessard 1991) and for ciliates fixed with 5% formalin (Putt & Stoecker 1989).

Microscopic analyses revealed that samples taken from replicate holes drilled 1 m apart at a site often differed considerably in the abundance and species composition of the phagotrophic protist assemblage. Variability among 'replicate' samples could be due either to counting errors or to fine-scale horizontal variability (differences among holes at a site). In order to evaluate these possibilities, samples were taken from 3 replicate holes spaced approximately 1 m apart and 3 replicate subsamples from each sample (i.e. hole) were analyzed microscopically. A random effects (Model II) 1-way analyses of variance was used to estimate the components of the variance associated with

small scale horizontal variability and counting error (Meyers & Milton 1991).

For scanning electron microscopy (SEM), samples were preserved in 10% Bouin's solution and settled on polylysine coated coverslips. Samples were then post-fixed in osmium tetroxide, serially dehydrated in acetone in 10% steps and critically point dried (Thomsen et al. 1991).

## RESULTS

### Physical and chemical parameters

At the ice edge sites, brine temperature ranged between  $-1.85$  and  $-0.3^\circ\text{C}$ , brine salinity ranged between 0 and 36 psu (practical salinity units) and chlorophyll *a* content varied between  $0.1$  and  $3.8 \mu\text{g l}^{-1}$  of brine (Table 1). Contrary to expectations, at H90 the land-fast ice decayed faster than at some more northern locations in McMurdo Sound. By early January, the sea-ice at H90 had become very porous (Stoecker et al. 1992), the brine temperature had risen to between  $-0.2$  and  $0^\circ\text{C}$ , and the brine salinity had decreased to  $\leq 2$  psu (Table 2).

Table 1. Mean densities (SD) of phagotrophic protists in the upper sea-ice brine, ice edge sites, McMurdo Sound, 1990–91. On each sampling date 2 to 4 samples were obtained ca 1 m apart at each site. The mean for each time period was calculated from the means for each date/site

	Dimensions ( $\mu\text{m}$ )	15–31 Dec	1–15 Jan	16–31 Jan
No. of sampling dates		4	4	3
Avg. no. of samples date <sup>-1</sup>		2.7	3.5	2.3
Brine temperature ( $^{\circ}\text{C}$ )		-1.85 to -0.3	-0.9 to -0.3	-1.5 to -0.5
Brine salinity (psu)		9 to 36	12 to 29	19 to 25
Chl a ( $\mu\text{g l}^{-1}$ )		0.2 to 3.8	0.1 to 0.9	0.1 to 0.9
Phagotrophic protists (cells $\text{l}^{-1}$ )				
Flagellates				
Unident. heterotrophic flag. $\leq 5 \mu\text{m}$	2 to 4	$3.2 \times 10^5$ ( $1.8 \times 10^5$ )	$1.9 \times 10^5$ ( $2.3 \times 10^5$ )	$0.8 \times 10^5$ ( $0.9 \times 10^5$ )
<i>Cryothecomomas</i>				
<i>C. armigera</i>	$7 \times 14$	$2.3 \times 10^4$ ( $2.6 \times 10^4$ )	$5.4 \times 10^4$ ( $9.3 \times 10^4$ )	< 50
<i>Cryothecomomas</i> sp.	$5 \times 7$	$1.3 \times 10^6$ ( $2.2 \times 10^6$ )	$1.0 \times 10^6$ ( $1.7 \times 10^6$ )	< 50
Heterotrophic dinoflagellates				
<i>Gymnodinium</i> sp.	$10 \times 17$	$0.3 \times 10^4$ ( $0.5 \times 10^4$ )	$2.2 \times 10^4$ ( $2.8 \times 10^4$ )	$1.9 \times 10^4$ ( $1.2 \times 10^4$ )
<i>Polykrikos</i> sp.	$34 \times 121$	45 (52)	114 (37)	35 (50)
<i>Gyrodinium</i> sp.	$34 \times 129$	2 (5)	22 (28)	19 (12)
<i>Protoperdinium</i> sp.	35 to 40	< 10	15 (26)	< 10
Ciliates (total) <sup>a</sup>				
Oligotrichs (total)		$1.2 \times 10^3$ ( $0.7 \times 10^3$ )	$4.8 \times 10^3$ ( $3.8 \times 10^3$ )	$3.6 \times 10^3$ ( $2.2 \times 10^3$ )
Selected species		$0.7 \times 10^3$ ( $0.5 \times 10^3$ )	$2.9 \times 10^3$ ( $2.0 \times 10^3$ )	$2.1 \times 10^3$ ( $0.4 \times 10^3$ )
<i>Strombidium</i> sp. 2 <sup>b</sup>	$25 \times 30$	46 (80)	71 (126)	471 (816)
<i>Strombidium</i> sp. 3	$20 \times 25$	110 (207)	185 (232)	72 (119)
<i>Strombidium</i> sp. 4	$45 \times 60$	204 (317)	1100 (700)	751 (1100)
<i>Strombidium</i> sp. 6 <sup>b</sup>	30	136 (220)	< 10	387 (671)
<i>Strombidium</i> sp. 9 <sup>b</sup>	$30 \times 60$	24 (48)	909 (1557)	243 (328)
<i>Strombidium</i> sp. 14	15 to 20	25(50)	264 (404)	19 (16)
<i>Mesodinium rubrum</i> <sup>b</sup>	20 to 30	156 (151)	1830 (1779)	728 (852)
<i>Didinium</i> spp.	25 to 60	< 10	< 10	740 (1000)
<i>Scuticociliate</i> sp.	$10 \times 25$	340 (630)	< 10	7 (12)

<sup>a</sup>Taxa with average densities  $> 10^2$  are listed individually as well as of part of the total; <sup>b</sup>plastidic species

### Composition and abundance of heterotrophic flagellates and ciliates

The phagotrophic protist assemblage of the upper land-fast ice brine community consisted of unidentified  $< 5 \mu\text{m}$  heterotrophic flagellates, flagellates in the genus *Cryothecomomas*, heterotrophic dinoflagellates and ciliates (Tables 1 & 2, Fig. 2). The  $< 5 \mu\text{m}$  category was comprised of heterotrophic cells about  $2 \times 4 \mu\text{m}$  in size; these generally occurred at densities of  $10^4$  to  $10^5$  cells  $\text{l}^{-1}$  and made up  $< 3\%$  of the heterotrophic protist biomass (Figs. 3 & 4).

Two species of *Cryothecomomas*, *C. armigera* Thomsen & Buck 1991 and an unidentified 5 to 7  $\mu\text{m}$  species were present in the brine at some sites (Fig. 2). During late December and early January, the average density of *C. armigera* at the ice edge sites was high,  $10^4$  to  $10^5$  cells  $\text{l}^{-1}$ , but it decreased to  $< 50$  cells  $\text{l}^{-1}$  by late January (Table 1). The temporal pattern of abundance of *Cryothecomomas* sp. at H90 was similar, except that

the peak and drop in abundance occurred earlier than at the ice edge sites (Table 2). Cells of *Mantoniella* sp. were observed within the food vacuoles of *C. armigera*. The smaller, unidentified *Cryothecomomas* sp. occurred at average densities of  $10^6$  cells  $\text{l}^{-1}$  in the samples collected near the ice edge during late December and early January (Table 1); it was not observed in the H90 samples. Although *Cryothecomomas* spp. occurred sporadically in samples (Tables 3 & 4), this genus contributed on average over 50% of the biomass of heterotrophic protists in the ice edge brine samples during the last half of December and in the H90 samples taken in late November (Figs. 3 & 4). Near the end of the sampling period, the contribution of *Cryothecomomas* to biomass in the brine was negligible (Figs. 3 & 4).

Heterotrophic dinoflagellates were a common component of the protist assemblage (Tables 1 & 2). On average, this group contributed 5 to 15% of the biomass of heterotrophic protists in the brine samples collected near the ice edge (Fig. 3). The dominant species

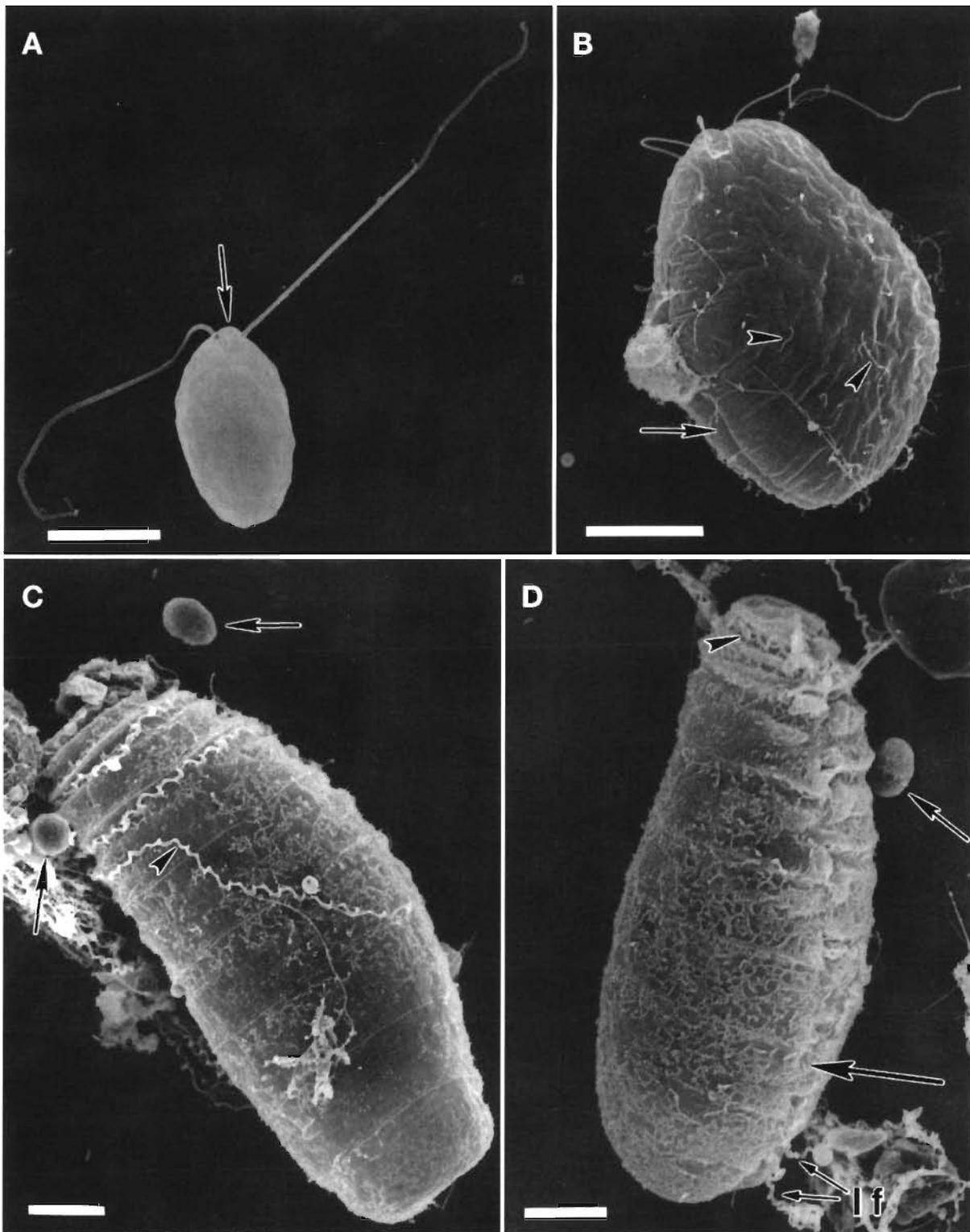


Fig. 2. Scanning electron micrographs. (A) Small *Cryothecomonas* sp. with 2 slightly unequal apically inserted flagella arising from either side of a papilla (arrow). Scale bar = 5  $\mu\text{m}$ . (B) *C. armigera* with 2 apically inserted flagella, cytostomal slit (arrow) and discharged extrusomes (arrowheads). Scale bar = 10  $\mu\text{m}$ . (C) *Polykrikos* sp. with cingulum and transverse flagellum. Small vegetative form of the dominant autotrophic dinoflagellate (arrows) that is a major component of the brine and which is preyed on by *Polykrikos* sp. Scale bar = 10  $\mu\text{m}$ . (D) *Polykrikos* sp. showing cingulum with transverse flagellum (arrowhead), sulcus (large arrow) and longitudinal flagella (lf). Stage had to be tilted 30° to obtain this picture of the sulcus. Small vegetative form of the dominant autotrophic dinoflagellate is indicated by a small arrow. Scale bar = 10  $\mu\text{m}$

Table 2. Changes (means and SD) in phagotrophic protist densities in the upper sea-ice brine, H90, McMurdo Sound, 1990–91. ND: no data

	30 Nov	12 Dec	24 Dec	5 Jan
Brine temp. (°C) <sup>a</sup>	-2.0	-1.0 to -1.3	-0.7 to -0.2	-0.2 to 0.0
Brine salinity (psu) <sup>a</sup>	51 to 52	31 to 42	3 to 31	<1 to 2
Chl a (µg l <sup>-1</sup> ) <sup>a</sup>	0.32 to 1.14	1.2 to 3.9	0.3 to 0.8	0.3 to 0.5
Protists (cells l <sup>-1</sup> ) <sup>b</sup>				
Flagellates				
Unidentified heterotrophic flagellates ≤ 5 µm	ND	5.7 × 10 <sup>4</sup> (6.4 × 10 <sup>4</sup> )	2.0 × 10 <sup>4</sup> (3.1 × 10 <sup>4</sup> )	1.4 × 10 <sup>4</sup> (2.0 × 10 <sup>4</sup> )
<i>Cryothecomonas armigera</i>	1.0 × 10 <sup>4</sup> (1.1 × 10 <sup>4</sup> )	9.7 × 10 <sup>4</sup> (2.0 × 10 <sup>4</sup> )	1.8 × 10 <sup>4</sup> (1.6 × 10 <sup>4</sup> )	<50
Heterotrophic dinoflagellates				
<i>Gymnodinium</i> sp.	<10	<10	<10	<10
<i>Polykrikos</i> sp.	7 (14)	<10	<10	<10
<i>Gyrodinium</i> sp.	<10	<10	<10	14 (28)
Ciliates (total)				
Oligotrichs	93 (106)	ND	200 (160)	7 (14)
<i>Mesodinium rubrum</i>	<10	ND	80 (69)	<10

<sup>a</sup>Data from Stoecker et al. (1992); <sup>b</sup>average of 3 replicates taken 1 m apart

was a small colorless *Gymnodinium* sp. about 10 × 17 µm in size; it occurred in about 52 % of the brine samples collected at the ice edge (Table 4). Its average density was between 10<sup>4</sup> and 10<sup>5</sup> cells l<sup>-1</sup> in ice edge

brine samples collected near the end of the sampling season (Table 1). On average, the *Gymnodinium* sp. contributed an estimated 64 % of the total heterotrophic dinoflagellate biomass.

Table 3. Fine-scale spatial variability (mean cells l<sup>-1</sup> with coefficient of variation, CV) in the brine assemblage at H90, 24 December 1991. Samples A, B and C were obtained from 0.5 m deep holes drilled in the sea-ice at 1 m intervals along a 3 m transect. Three subsamples from each hole were enumerated in order to determine the within-sample variation (counting error). For *Cryothecomonas* spp., 75 % of the total variation was due to differences among holes. For total ciliates, 91 % of the total variation was due to differences among holes

<b><i>Cryothecomonas</i> spp.</b>				
	Hole A	Hole B	Hole C	
Density	28.9 × 10 <sup>3</sup> (16.3 %)	25.1 × 10 <sup>3</sup> (35.0 %)	0 (0 %)	
Source of variation	df	MS	F	
Among different holes	2	739.8	10.2116, p < 0.05	
Within the same hole	6	72.4		
<b>Total ciliates</b>				
	Hole A	Hole B	Hole C	
Density	567 (15.9 %)	213 (19.5 %)	36.7 (68.6 %)	
Source of variation	df	MS	F	
Among different holes	2	218477.8	32.9916, p < 0.001	
Within the same hole	6	6622.2		

Three > 20 µm in size heterotrophic dinoflagellates were observed, *Polykrikos* sp., *Gyrodinium* sp. and a small *Proto-peridinium* sp. (Tables 1 & 2). The *Polykrikos* sp. was over 100 µm long (Fig. 2) and microscopic evidence indicates that it preys on the ca 10 µm photosynthetic dinoflagellate that was the dominant alga in the brine (Stoecker et al. 1992). *Polykrikos* sp. occurred in about 67 % of the samples (Table 4), had average densities between 35 and 114 cells l<sup>-1</sup> (Table 1), and accounted for an estimated 28 % of the heterotrophic dinoflagellate biomass in brine samples collected near the ice edge. *Gyrodinium* sp. and *Proto-peridinium* sp. occurred sporadically in the brine (Table 4).

The ciliate assemblage in the sea-ice brine was diverse; it included members of the orders Choreotrichida (e.g. tintinnids and *Strobilidium*), Oligotrichida (e.g. *Strombidium* and *Laboea*), Prorodontida (e.g. *Spiroprorodon*), Haptorida (e.g. *Didinium*, *Lacrymaria*, *Mesodinium*), Euplotida (e.g. *Aspidisca*, *Euplotes*), Scuticociliatida and other ciliates of unknown affinities. Except at H90, ciliate densities averaged between 1.0 and 5.0 × 10<sup>3</sup> cells l<sup>-1</sup> (Table 1) and cili-

Table 4. Frequency of occurrence (no. of samples in which a species was present out of 3 samples taken at 1 m intervals) in the upper sea-ice brine, ice edge sites, McMurdo Sound, 1990–1991. Data only from sites/dates from which  $\geq 3$  samples were taken. If more than 3 samples were taken, 3 were randomly chosen for analyses here; thus not all species in Table 1 have a percent occurrence  $> 0$  in this table. Data are only presented for species with average densities  $> 10^2$  cells  $l^{-1}$

	Dec		Jan				Percent of samples present	
	26	29	5	7	10	21		23
<b>Heterotrophic flagellates</b>								
<i>Cryothecomonas</i>								
<i>C. armigera</i>	3	0	0	0	0	1	0	19%
<i>Cryothecomonas</i> sp.	0	0	0	0	0	0	0	0%
<b>Heterotrophic dinoflagellates</b>								
<i>Gymnodinium</i> sp.	0	2	1	3	1	3	1	52%
<i>Polykrikos</i> sp.	1	1	3	3	2	2	2	67%
<i>Gyrodinium</i> sp.	0	0	0	2	0	0	0	10%
<i>Protoperdinium</i> sp.	0	0	0	1	0	0	0	5%
<b>Ciliates</b>								
<i>Strombidium</i> sp. 2	1	2	1	1	0	3	0	38%
<i>Strombidium</i> sp. 3	1	0	0	1	2	1	0	24%
<i>Strombidium</i> sp. 4	0	2	3	3	3	3	2	76%
<i>Strombidium</i> sp. 6	0	0	0	0	0	0	1	5%
<i>Strombidium</i> sp. 9	0	2	0	3	0	2	1	38%
<i>Strombidium</i> sp. 14	0	1	1	2	1	3	1	43%
<i>Mesodinium rubrum</i>	1	3	3	3	3	3	3	90%
<i>Didinium</i> sp.	0	0	2	1	0	0	3	28%
<i>Scuticociliate</i> sp.	1	3	0	0	0	0	0	19%

ates comprised on average 15 % in late December and 48 to 85 % in January of the heterotrophic protist biomass (Fig. 3).

The most common ciliates in the brine were oligotrichs; they comprised about 70 % of the ciliate biomass (data not shown). Six *Strombidium* spp. dominated the oligotrich assemblage and occurred at average individ-

ual species densities of  $\geq 100$   $l^{-1}$  of brine (Table 1). The most common oligotrich was *Strombidium* sp. 4, a large cone-shaped species (ca  $45 \times 60$   $\mu m$ ) with a prominent sheath capping its posterior end; microscopic evidence indicated that this species primarily preyed on the 10  $\mu m$  photosynthetic dinoflagellate found in the brine (Stoecker et al. 1992). This *Strombidium* sp. occurred at average densities of ca  $10^3$  cells  $l^{-1}$  and in 76 % of the samples taken near the ice edge (Tables 1 & 4). It accounted for about 45 % of the total ciliate biomass (data not shown). Three plastidic *Strombidium* species, (spp. 2, 6 and 9) occurred at average combined densities of ca  $10^3$   $l^{-1}$  of brine at the ice edge sites (Table 1; Stoecker et al. 1992).

A tintinnid, *Laackmanniella* sp. (oral diameter 25 to 35  $\mu m$ , length 75 to 150  $\mu m$ ), only occurred in the 4 brine samples collected near the ice edge on 7 January. On this 1 date/location, it occurred at an average density of 226 (SD 197)  $l^{-1}$ .

The ciliate which occurred most frequently in the brine was the autotrophic species *Mesodinium rubrum* (= *Myrionecta rubra*); it occurred in 90 % of the samples collected near the ice edge (Table 4) and was also observed at H90 (Table 2). Average *M. rubrum* densities were 156 cells  $l^{-1}$  in the latter part of December, 1830 cells  $l^{-1}$  in the beginning of January, and 728 cells  $l^{-1}$  at the end of January in the brine samples collected near the ice edge (Table 1). On average, *M. rubrum* accounted for about 14 % of the ciliate biomass (data not shown).

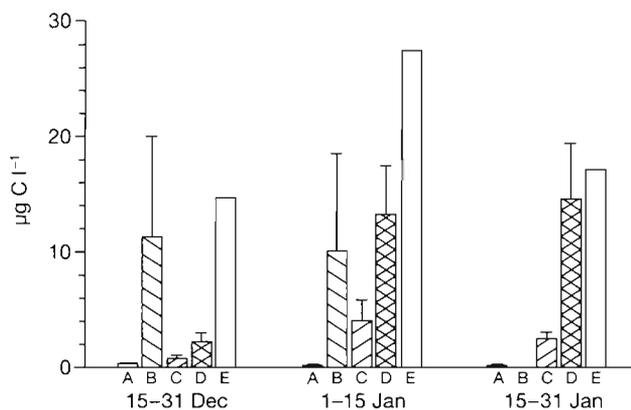


Fig. 3. Changes (+ SD) in the biomass of (A)  $\leq 5$   $\mu m$  heterotrophic flagellates, (B) *Cryothecomonas* spp., (C) heterotrophic dinoflagellates, (D) ciliates and (E) total protozoan biomass in the upper sea-ice brine at sites near the ice edge

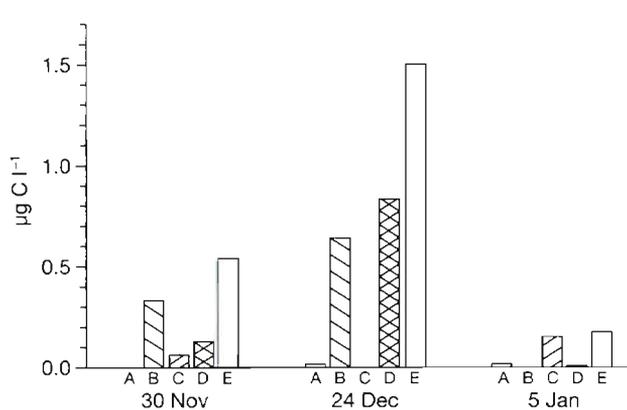


Fig. 4. Changes in the biomass of (A)  $\leq 5$   $\mu m$  heterotrophic flagellates, (B) *Cryothecomonas* spp., (C) heterotrophic dinoflagellates, (D) ciliates and (E) total protozoan biomass in the upper sea-ice brine at H90

The predatory ciliate *Didinium* spp. was not an important member of the assemblage until late in the season (Table 1). It occurred in 28% of the brine samples collected near the ice edge and had an average density of 740 cells  $l^{-1}$  in late January (Tables 1 & 4). *Spiroprorodon*, *Aspidisca* and *Euplotes* spp. did not appear in the brine until the latter half of January, providing a minor (<10%) contribution to ciliate biomass (data not shown). However, it is likely that ciliates that are associated with surfaces, such as *Aspidisca* and *Euplotes*, were underestimated by our sampling technique.

An unidentified scuticociliate ca  $10 \times 20$  to  $25 \mu m$  in size occurred in 19% of the samples collected near the ice edge and in the latter half of December occurred at average densities of 340 cells  $l^{-1}$  although it was rarer in January (Tables 1 & 4). Scuticociliates are often associated with surfaces and may also have been underestimated by our sampling technique. The scuticociliates contributed <1% of the total ciliate biomass in our samples (data not shown). Photosynthetic prey were not visible within this species.

#### Horizontal spatial variability in the brine assemblage

For all heterotrophic protists encountered in the brine, variation in density was high (Tables 1 & 2). Among the  $>5 \mu m$  species that occurred at average densities  $>10^2$  cells  $l^{-1}$ , most occurred in <50% of the samples (Table 4). Variability due to counting technique could potentially cause this pattern. Thus, we compared variability among holes and among replicate counts using a random effects ANOVA. For 2 taxa analyzed in this fashion, we estimated that between 75 and 91% of the total variability in cell densities was due to differences among holes (Table 3). Differences among samples taken on the same date at the same site were primarily due to fine scale horizontal spatial variability rather than to counting errors. (Our sampling method did not permit analyses of vertical spatial variability).

#### Temporal patterns in the contribution of photosynthetic and non-photosynthetic protists to biomass

The contribution of heterotrophs to the protist biomass in the upper land-fast ice brine increased from ca 10% in December to an average of 57% during the latter half of January in the samples taken near the ice edge (Table 5). In the samples from H90, the heterotrophic component of the protistan biomass was <10% and appeared to peak in late December (Table 5).

Table 5. Comparison of the average biomass ( $\mu g C l^{-1}$ ) of photosynthetic and non-photosynthetic protists in the upper sea-ice brine, McMurdo Sound, 1990–91

	Photo-synthetic <sup>a</sup>	Non-photosynthetic	Percent strictly heterotrophic
Ice edge sites			
15–31 Dec	129.0	14.7	10.2%
1–15 Jan	79.1	27.6	25.9%
15–31 Jan	13.1	17.1	56.6%
H90			
30 Nov	235.2	0.5	0.2%
24 Dec	20.1	1.5	7.0%
5 Jan	34.5	0.2	0.5%

<sup>a</sup> Data from Stoecker et al. (1992); includes photosynthetic mixotrophic ciliates

## DISCUSSION

The microalgal assemblage in the upper land-fast ice brine of McMurdo Sound shows many similarities to the surface-layer and internal ice assemblages reported from pack-ice (Stoecker et al. 1992). Likewise, the heterotrophic assemblage is composed of taxa reported from pack-ice (Bartsch 1989, reviewed in Garrison 1991). *Cryothecomonas* spp. can be a dominant component of the biomass in both the land-fast ice brine (Figs. 2 & 3) and in pack-ice (Garrison & Buck 1991, Thomsen et al. 1991). *C. armigera* in both pack and land-fast ice appears to prey primarily on *Mantoniella* sp. and other  $<5 \mu m$  flagellates.

Heterotrophic dinoflagellates are present in both land-fast ice brine and surface-layer pack-ice communities (Bartsch 1989, Garrison & Buck 1989). However, we did not observe the large, spherical, athecate dinoflagellate that consumes diatoms and produces fecal pellets which has been observed in the surface slush layer of pack-ice (Buck et al. 1990). This difference may be due to the abundance of pennate diatoms in pack-ice and their scarcity in the land-fast ice brine (Garrison & Buck 1989, Stoecker et al. 1992).

The ciliate taxa that we observed in the land-fast ice brine have been reported from pack-ice (Fenchel & Lee 1972, Corliss & Snyder 1986, Garrison & Buck 1989). In both assemblages, ciliates can comprise a major fraction of the heterotrophic protist biomass, and oligotrichs (primarily *Strombidium* spp.) dominate the ciliate assemblage (Figs. 3 & 4, Tables 1 & 2; reviewed in Garrison 1991). In both pack-ice and land-fast ice brine, the autotrophic ciliate *Mesodinium rubrum* (*Myrionecta rubra*) is one of the most frequently observed species (Tables 1, 2 & 4; Garrison & Buck 1989).

Some taxa reported from pack-ice were not observed in our samples from the upper sea-ice brine, these included choanoflagellates and foraminifera (reviewed

in Garrison 1991). Foraminifera are known to be rare in congelation ice (Spindler et al. 1990) and land-fast ice is usually dominated by congelation ice (Garrison et al. 1986). The techniques used in the present study were not appropriate for counting amoebae or for identifying various taxa of  $<5 \mu\text{m}$  heterotrophic flagellates. Thus, many of the  $<5 \mu\text{m}$  species reported from land-fast ice in East Antarctica (Takahashi 1987) or from pack-ice (reviewed in Garrison 1991) may have been present but unreported in our samples as specific taxa. It is interesting to note that although micrometazoa are found in pack-ice and land-fast ice bottom type microbial communities (reviewed in Garrison 1991), we did not observe them in the upper land-fast ice brine. The grazers and predators in the upper land-fast ice brine were all protists.

Most of the heterotrophic flagellates and ciliates observed in pack-ice or land-fast ice are also found in the plankton (Garrison 1991, Thomsen et al. 1991). Internal communities in pack-ice are thought to be initiated from cells or cysts harvested from the water column by frazil ice or entrapped when the ice is formed (reviewed in Garrison 1991). It is likely that many of the species found in land-fast ice arrive in the same manner. However, it seems possible that some of larger protists found in the brine pockets, especially those only observed in the late austral spring or summer, may invade the ice from the water column.

It is interesting that many of the larger species most frequently found in the brine tend to swim toward light or upward. *Mesodinium rubrum* and *Strombidium* spp. can swim at speeds of up to  $1 \text{ m h}^{-1}$  or more (Crawford 1989, Jonsson 1989). Swimming speeds for heterotrophic dinoflagellates have not been reported, but photosynthetic ones have maximum swimming speeds of up to  $1 \text{ m h}^{-1}$  (Taylor & Pollinger 1987). Relatively long, straight brine channels are found in columnar sea-ice and in the late austral spring, when brine drainage occurs, the ice becomes extremely porous and mm to cm wide channels may be present at times (Maykut 1985, Eicken 1992, Weissenberger et al. 1992). Cysts of brine microalgae are found in the water column when brine drainage occurs (Stoecker et al. 1992, Stoecker unpubl. data) and therefore it seems possible that planktonic organisms may invade the ice at the same time.

A conspicuous feature of the heterotrophic protist assemblage is its fine-scale (order of meters) spatial variability in composition (Tables 3 & 4). We believe that this fine-textured spatial heterogeneity is due to the structure of land-fast sea-ice. In McMurdo Sound, the upper 1 m of ice cores is dominated by congelation (columnar) ice (Jeffries & Weeks 1991). In congelation sea-ice, brine channels are orientated vertically and

are long and straight (Maykut 1985, Weissenberger et al. 1992). Particularly during early and mid-spring, when brine volumes in the upper ice in McMurdo Sound are low (Buckley & Trodahl 1990), this vertical structure should limit horizontal dispersion of protists in the ice. When brine channels open to the underlying water column, the columnar structure should favor vertical migration of motile protists (such as planktonic ciliates and dinoflagellates) into the ice, where they may colonize brine pockets, but the columnar structure may still inhibit horizontal dispersion of the colonists.

Investigations of pack-ice microbial communities have suggested a successional sequence, with the heterotrophic components increasing in the spring and summer (Garrison & Buck 1989). As in the pack-ice (reviewed in Garrison 1991), land-fast ice protist populations are highly variable, but general successional trends are evident (Tables 1 & 2, Figs. 3 & 4). Average microalgal biomass  $\text{l}^{-1}$  of brine tends to decrease as heterotrophic protist biomass increases during the spring and early summer (Table 5). At H90, succession was truncated because of early decay of the ice. By 5 January at this site, the brine salinity was  $\leq 2$  psu and the upper ice was extremely porous (Table 2; Stoecker et al. 1992). Just before the ice decays (late January & February at most sites in McMurdo Sound), both protozoan and microalgal populations decrease (Table 5). At the ice edge sites, the average percent contribution of heterotrophs to the total protist biomass increased from ca 10% in late December to over 50% by the end of January (Table 5).

Changes in microalgal biomass  $\text{l}^{-1}$  of brine are partly due to dilution of the populations as brine volume increases, losses due to brine drainage and possibly due to death due to salinity changes, and encystment of the dominant species (Stoecker et al. 1992), but grazing is probably also important. The dominant phagotrophs, *Cryothecomonas* spp., the heterotrophic dinoflagellates and *Strombidium* spp., are all consumers of microalgae and other small eucaryotic cells.

The changes in heterotrophic protist biomass are probably due to a combination of growth of *in situ* populations, migration from the water column, dilution due to increases in brine volume and losses due to brine drainage, possibly death due to salinity changes, and predation. Some of the protists in the brine are predators on other heterotrophic protists. The small gymnodinoid dinoflagellate did not appear to contain plastidic food, it is likely that this species was a predator on *Cryothecomonas* spp. and other heterotrophic flagellates. At the ice edge, *Didinium* spp., which are usually voracious predators on other ciliates, were an important component of the brine during the end of January (Table 1).

Sea-ice, particularly multi-year and pack-ice, is notoriously heterogenous in its structure, nutrient content and in the distribution and abundance of microbial communities (reviewed in Garrison 1991). We hypothesize that the activities of phagotrophic protists are at least partly responsible for the observed heterogeneity in older sea-ice. The large, but variable, contribution of heterotrophs to land-fast ice brine protist assemblages just prior to ice decay and break-out suggests that grazing directly influences the distribution and abundance of bacteria and microalgae in ice floes that are formed from land-fast ice. Phagotrophic protists are thought to be the major contributors to nutrient regeneration in planktonic environments (Caron 1991), and it is likely that they play similar roles in brine communities. In pack-ice floes >2 mo in age, nitrate is often depleted and ammonia levels are often elevated (Garrison et al. 1990, Dieckmann et al. 1991). Levels and ratios of inorganic nutrients, as well as the composition of existing assemblages, must have an important influence on the development and species composition of ice microalgal communities in multi-year ice.

In Antarctica, pack-ice is more extensive than land-fast ice, but logistical problems make it difficult to document or experimentally investigate seasonal changes in the microbial communities of pack-ice. However, evidence strongly suggests that the activities of microorganisms (bacteria, microalgae and heterotrophic protists) influence ice decay and break-up (reviewed in Eicken et al. 1991) as well as being important in primary production in polar waters (reviewed in Legendre et al. 1992). Percent ice cover has an important influence on the structure and function of polar ecosystems as well as being an important parameter in global energy balance and atmosphere-ocean interactions (reviewed in Eicken 1992). Investigations of the microbial communities of sea-ice, and their interactions with their physical environment and the underlying water column, are necessary in order to understand polar ecosystem dynamics. The existence of upper sea-ice brine communities in land-fast ice, which is more accessible than pack-ice, should facilitate time-series studies and experimental investigations of the interactions between sea-ice microbial communities and their environment.

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